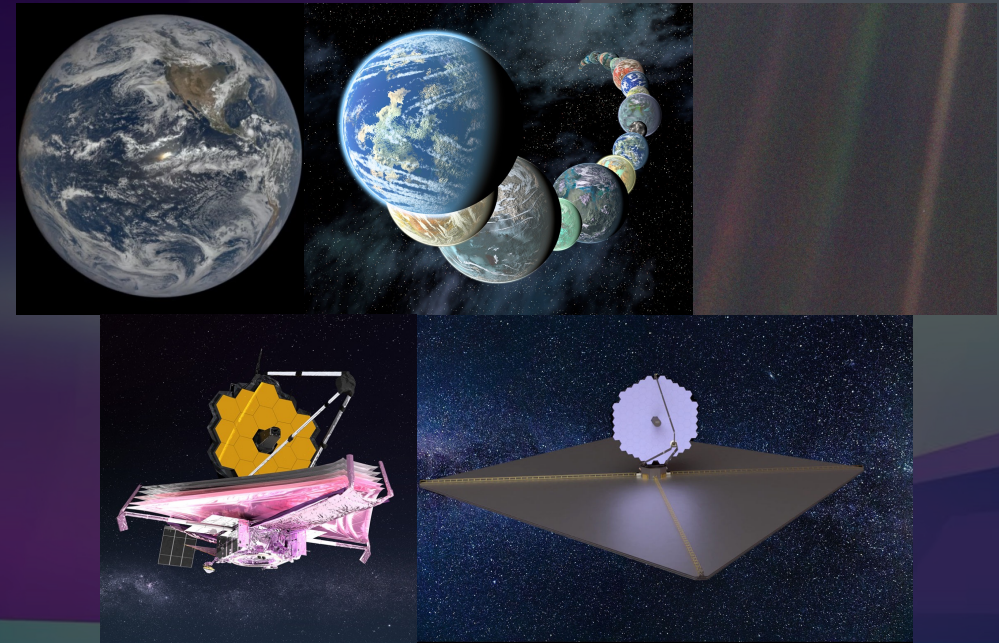
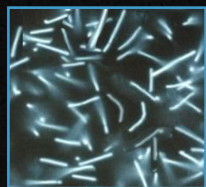
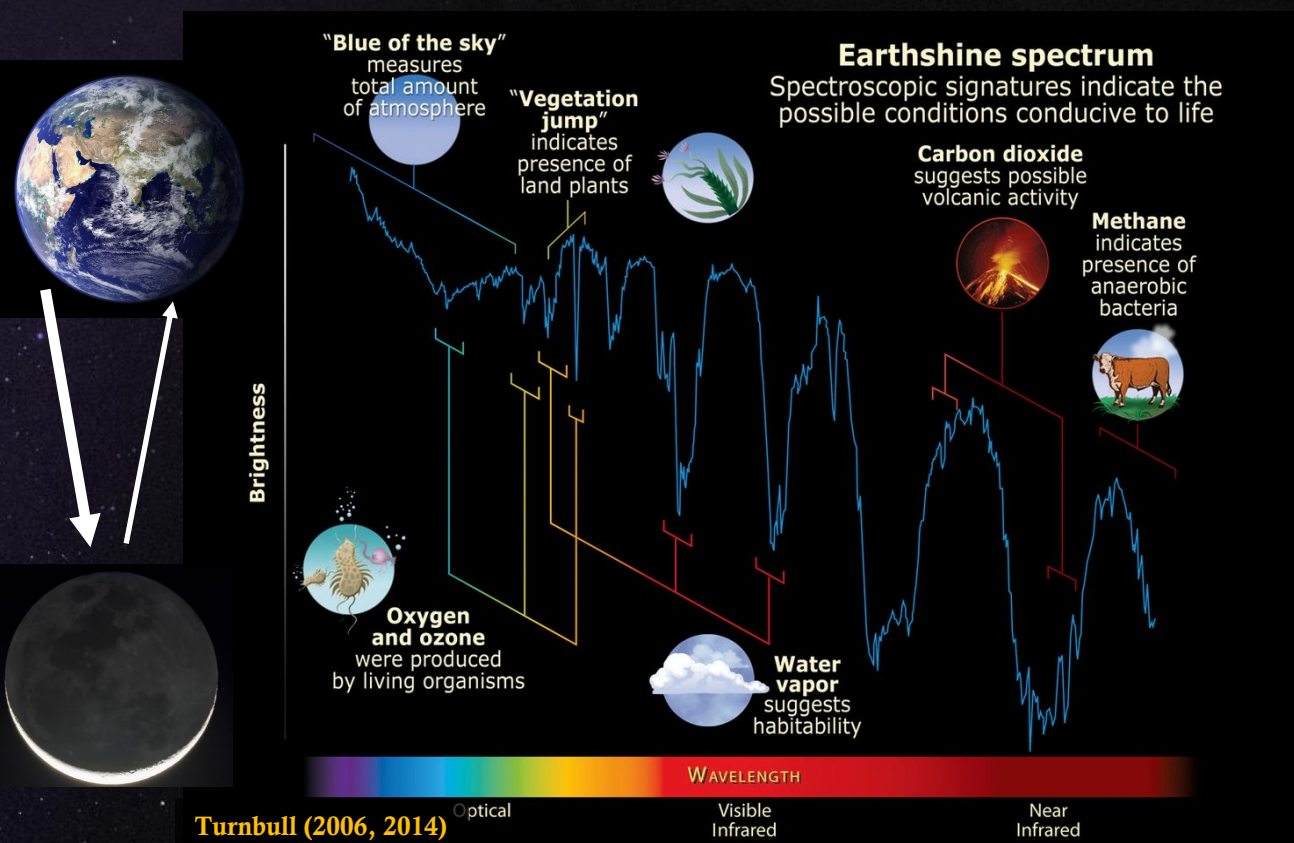


Earth as an Exoplanet: Relevance to ExoEarths

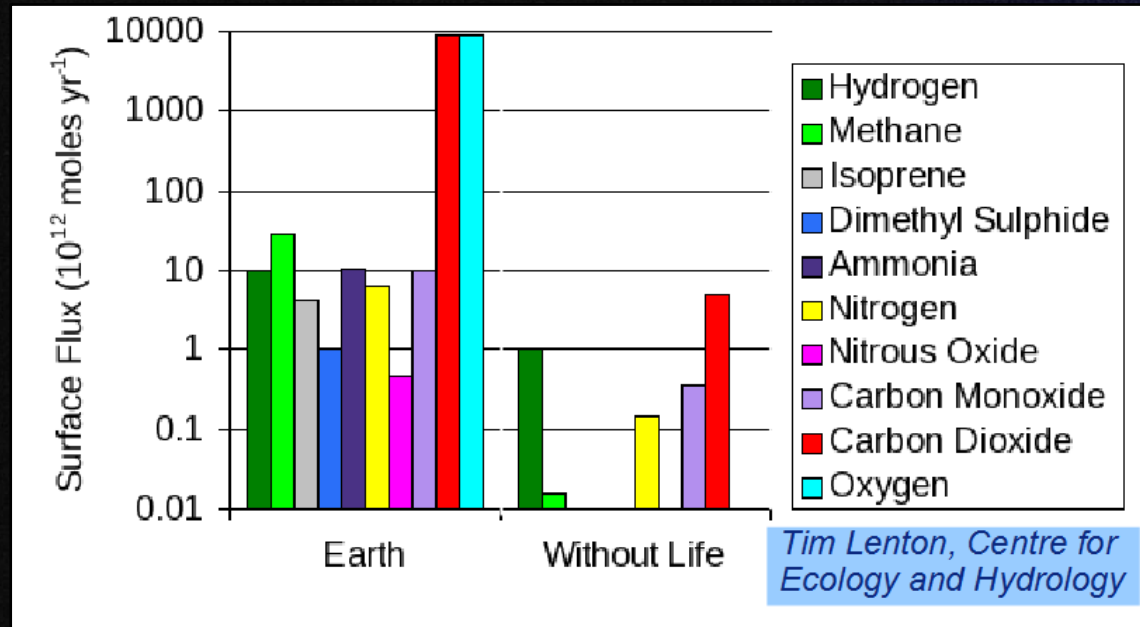
Edward Schwieterman
University of California, Riverside
ExoPAG 28
San Antonio, TX
10/1/2023



Earth is a Dynamic, Habitable, Geological Active, Living Planet

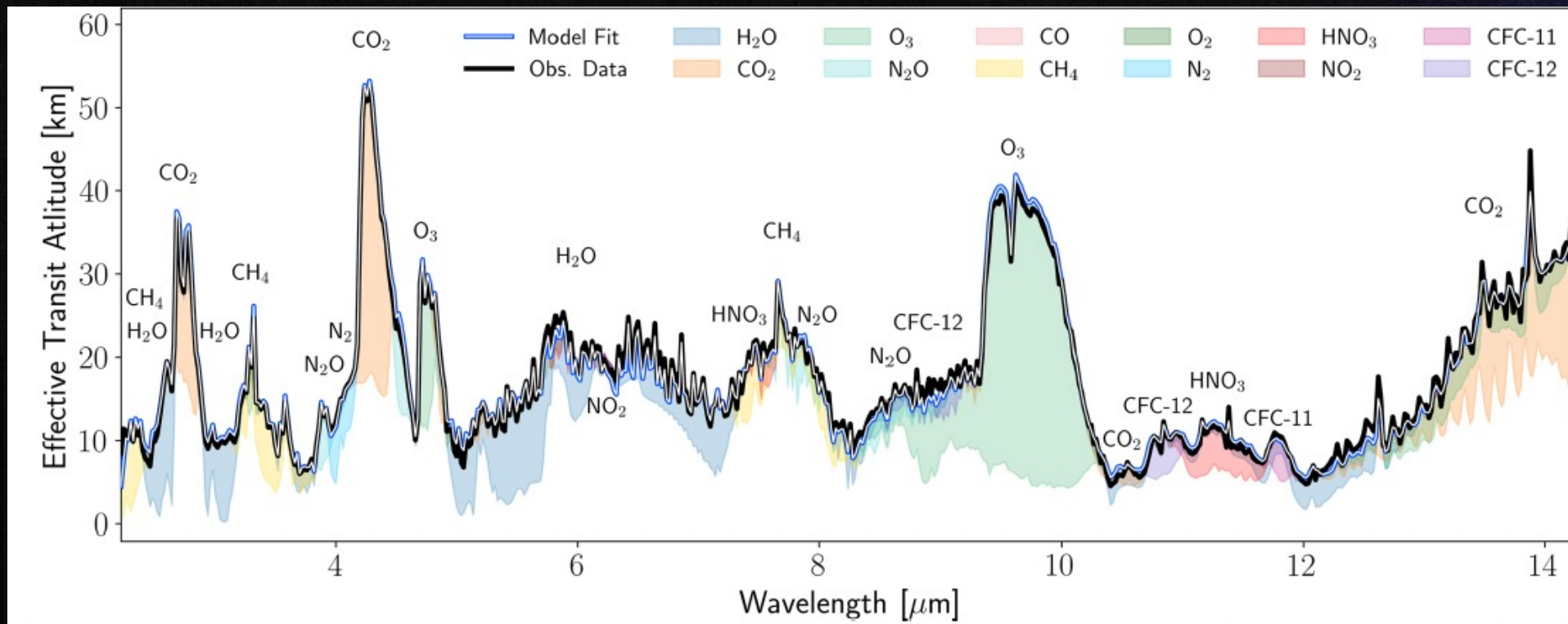
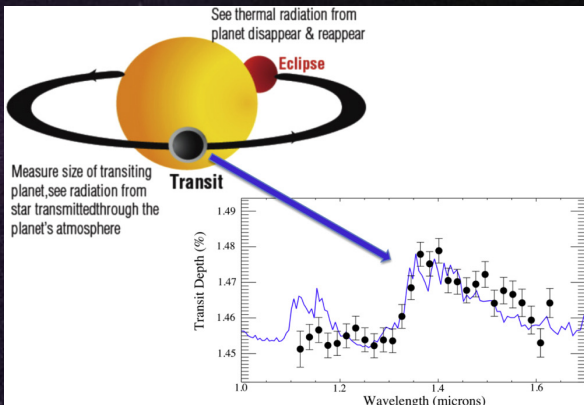
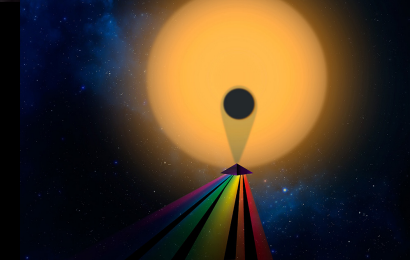


Fluxes of biogenic gases distinguish Earth from an abiotic counterpart



Life processes (e.g., carbon cycle, nitrogen cycle, oxygenic photosynthesis) have a large lever arm on Earth's atmospheric composition on geologically short timescales. These processes impact the remote spectrum.

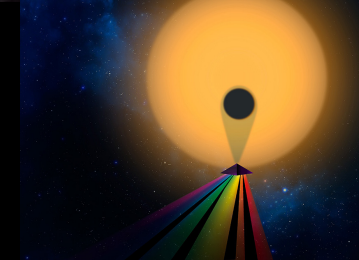
Earth's Transmitted Light Spectrum & Its Features



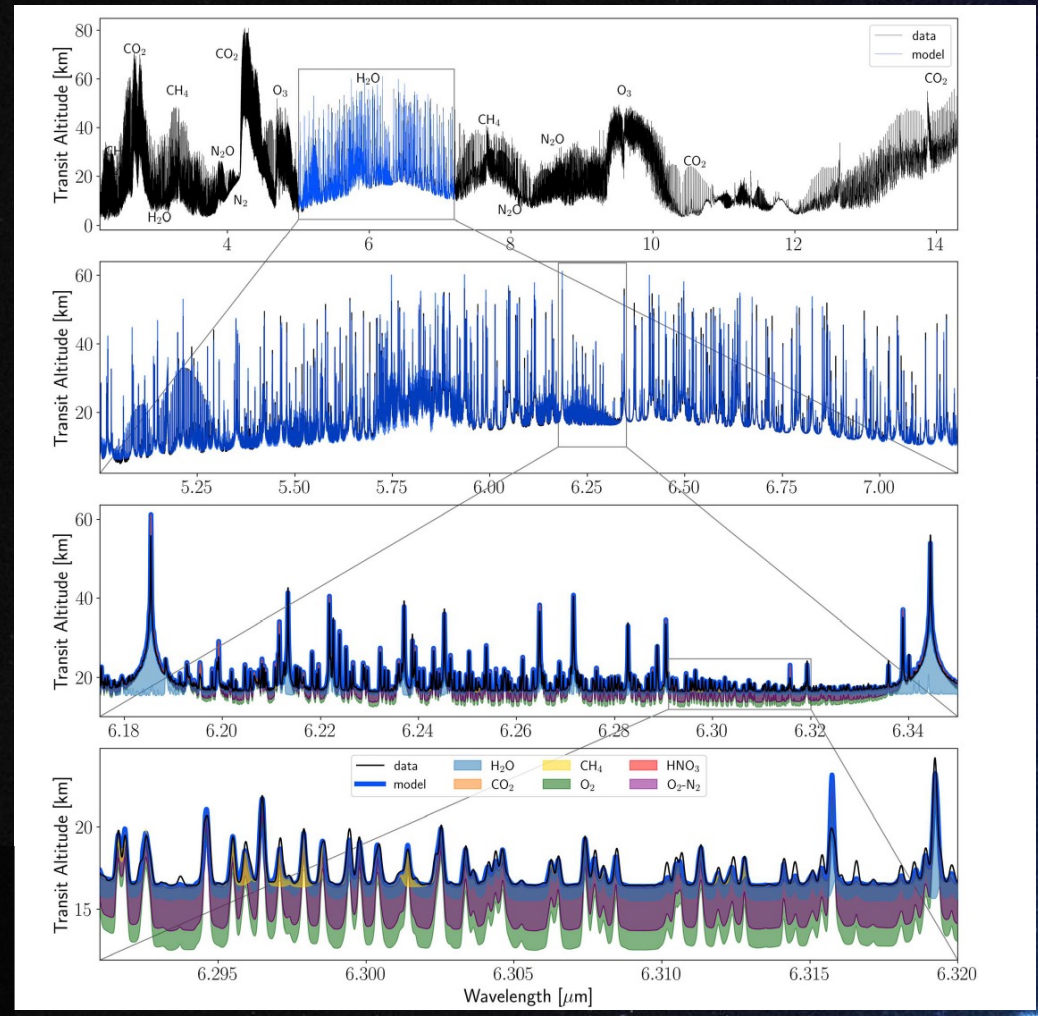
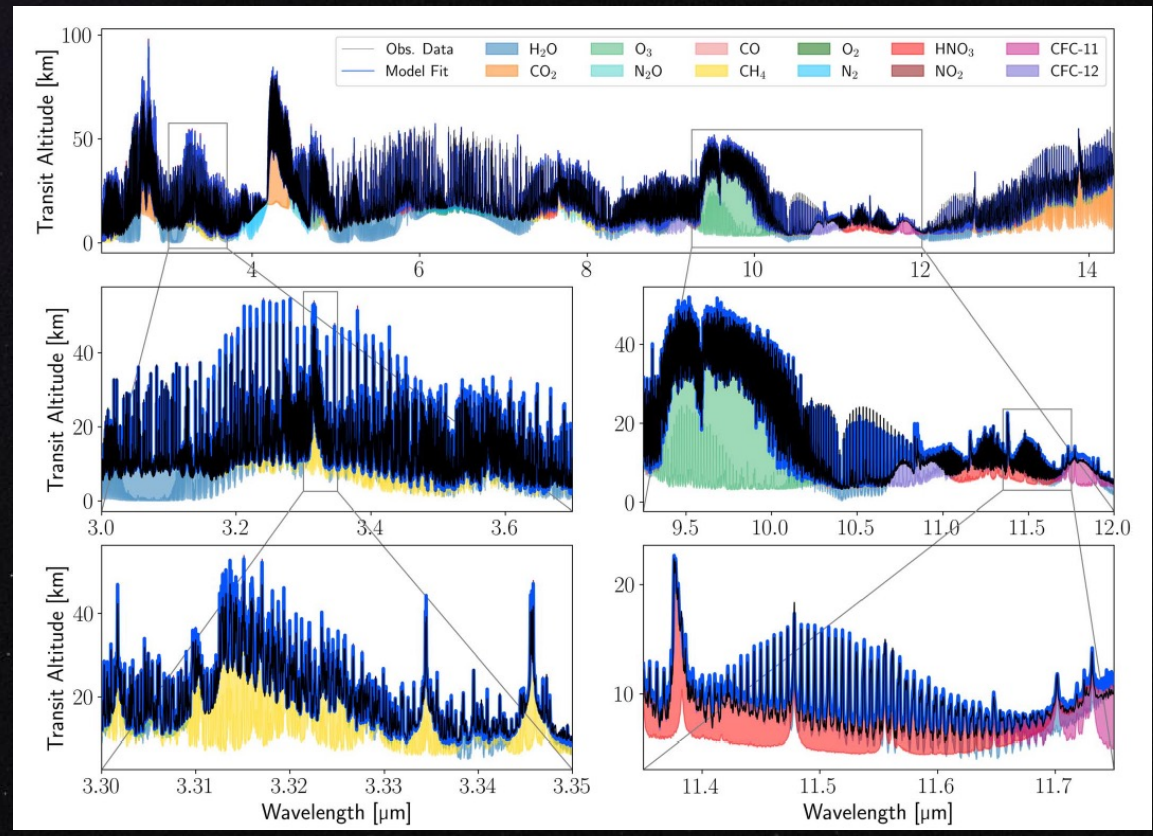
Top: illustration from Deming+2018; **Right:** model retrieval from Lustig-Yaeger+2023 using reconstructed Earth transmission spectrum from Macdonald & Cowan 2019, derived from SCISAT ACE-FTS data

Spectrum contains gaseous habitability markers **CO₂** (2.7, 4.3, 15 μm), **H₂O** (6 μm), and **N₂** (4.2 μm); biosignatures **CH₄** (3.3, 7.7 μm), **O₃** (4.7, 9.7 μm) and **N₂O** (4, 7.7, 8.5 μm) and technosignatures **CFC-11** (CCl₃F; 11.8 μm), **CFC-12** (CCl₂F₂), and **NO₂** (6.2 μm).

Earth's Transmitted Light Spectrum—High Resolution

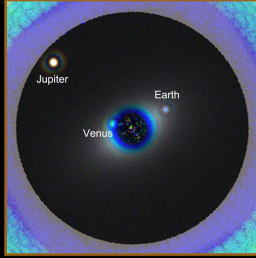


Transit Altitude [Km]

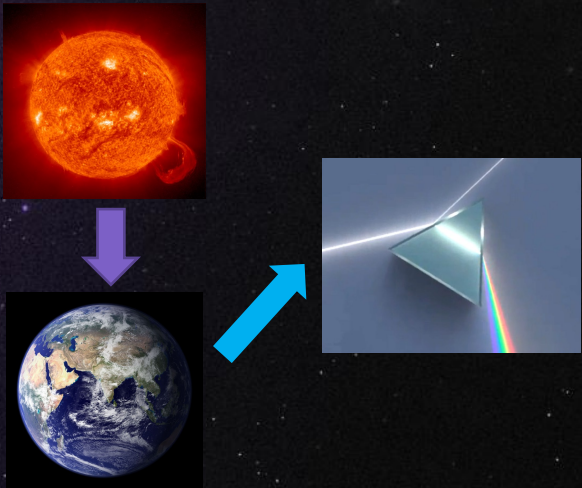


High-res data ($\Delta\nu = 0.02 \text{ cm}^{-1}$; $R=10^5$ at $5 \mu\text{m}$) are excellent for teasing out trace gas features and contributions from CIA ($\text{O}_2\text{-O}_2$; $\text{O}_2\text{-N}_2$)
 Macdonald & Cowan (2019, 2023); Lustig-Yaeger+2023

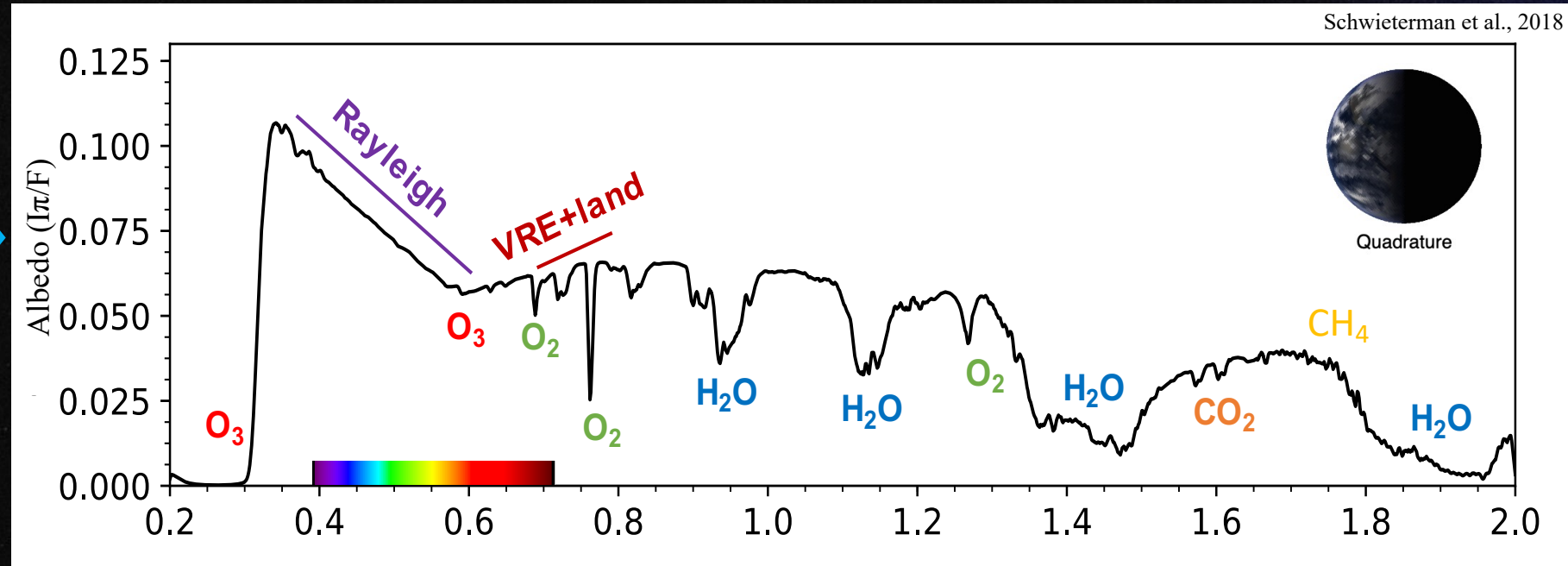
Earth's Reflected Light Spectrum and Its Features



Accessible Wavelength Regions will Depend on Observing Mode and Instrumentation

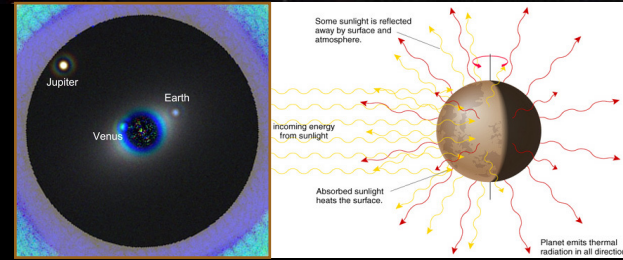


“Reflected light” – the light from the star is reflected (and/or scattered) by the planet to the distant observer.

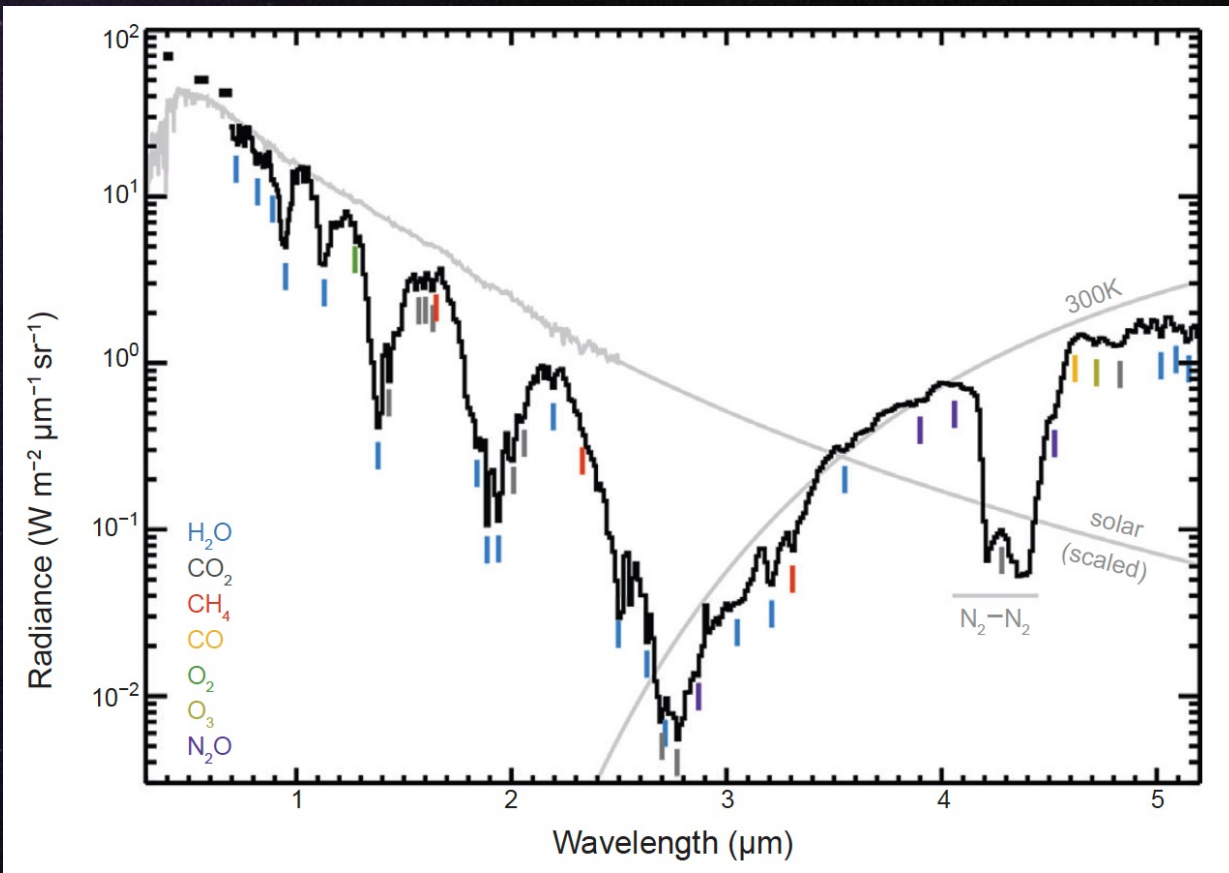


An Earth spectrum at quadrature (half illumination) in ultraviolet (UV), visible (VIS), and near-infrared (NIR) wavelengths shows features from oxygen (O_2), ozone (O_3), water vapor (H_2O), and carbon dioxide (CO_2). Rayleigh scattering from Earth's blue sky is shown (λ^{-4}). A small signature from Earth's vegetation red-edge (VRE) is apparent.

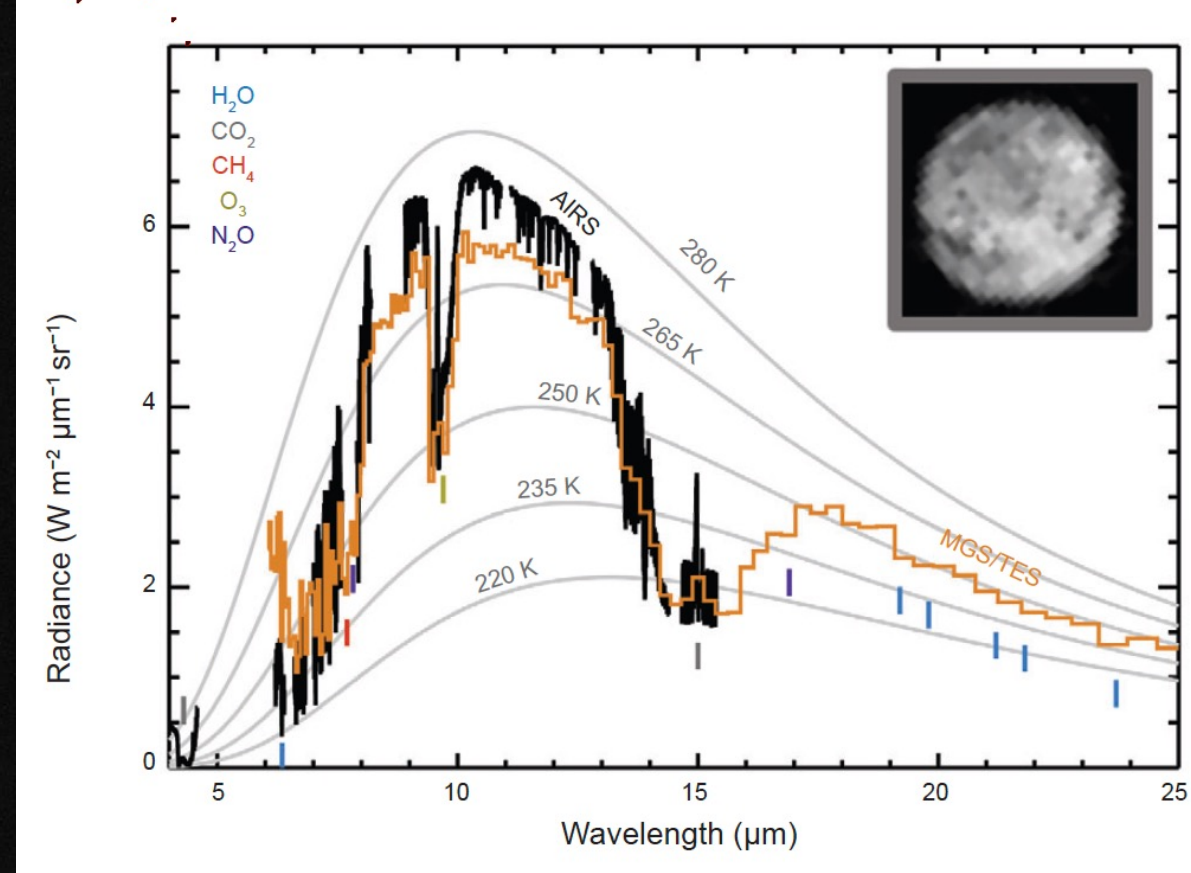
Importance of Wavelength-Range & Observing Mode



Earth's Reflected and NIR-emitted Light



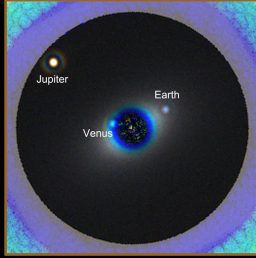
Earth's Thermal Infrared Spectrum



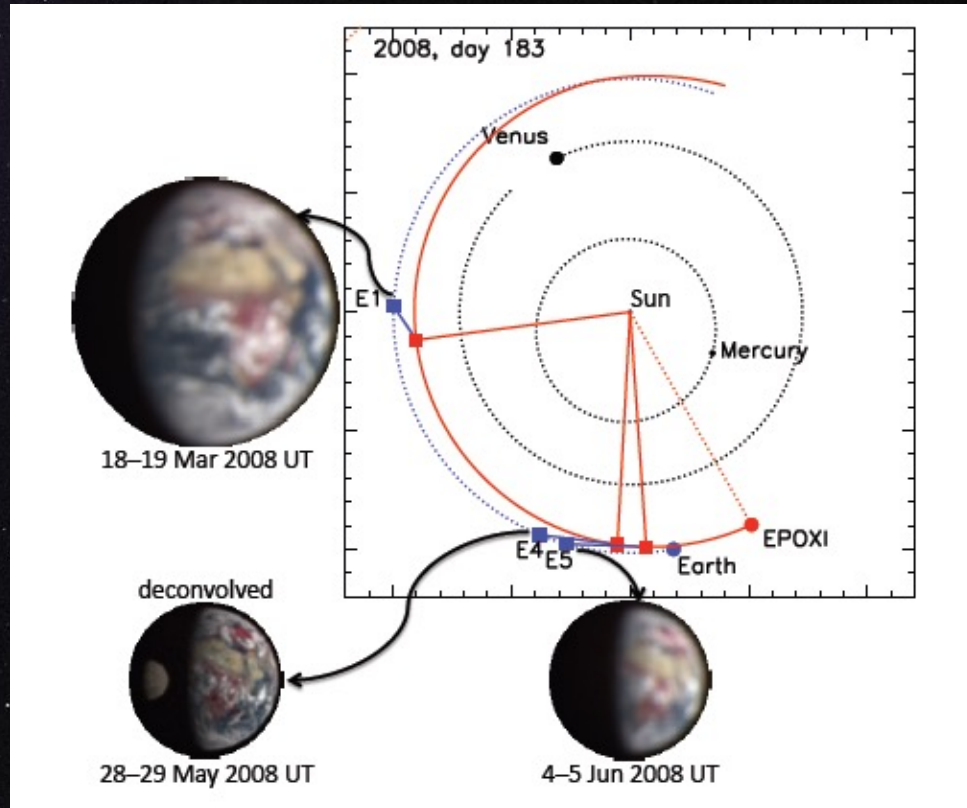
Interesting features where there is little flux...

Robinson & Reinhard, *Planetary Astrobiology*

Limited Existing Disk-Averaged, Phase-Dependent Data



EPOXI/Deep Impact (Livengood+ 2011)



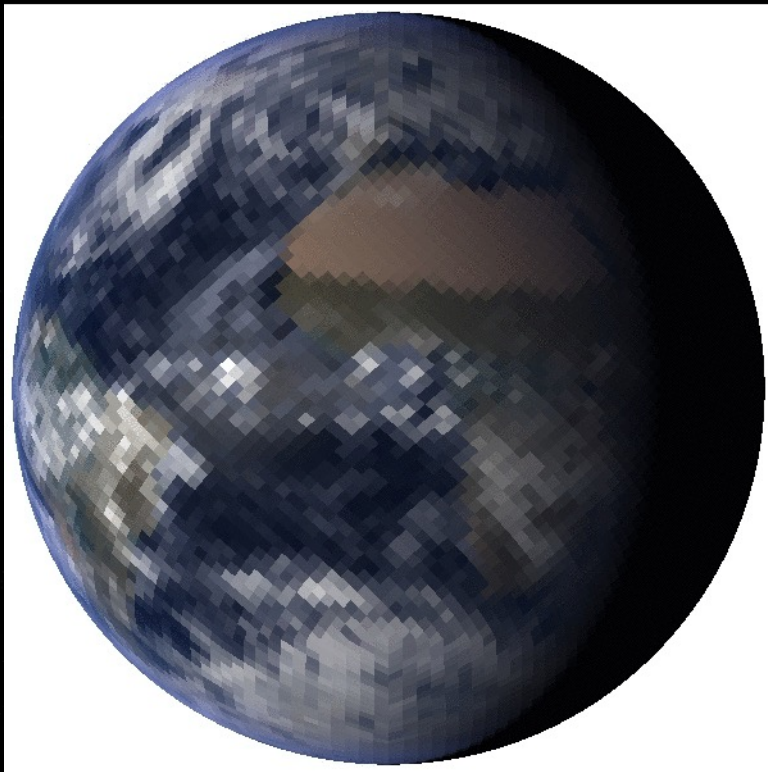
DSCOVR/EPIC (e.g, Carlson+2019)



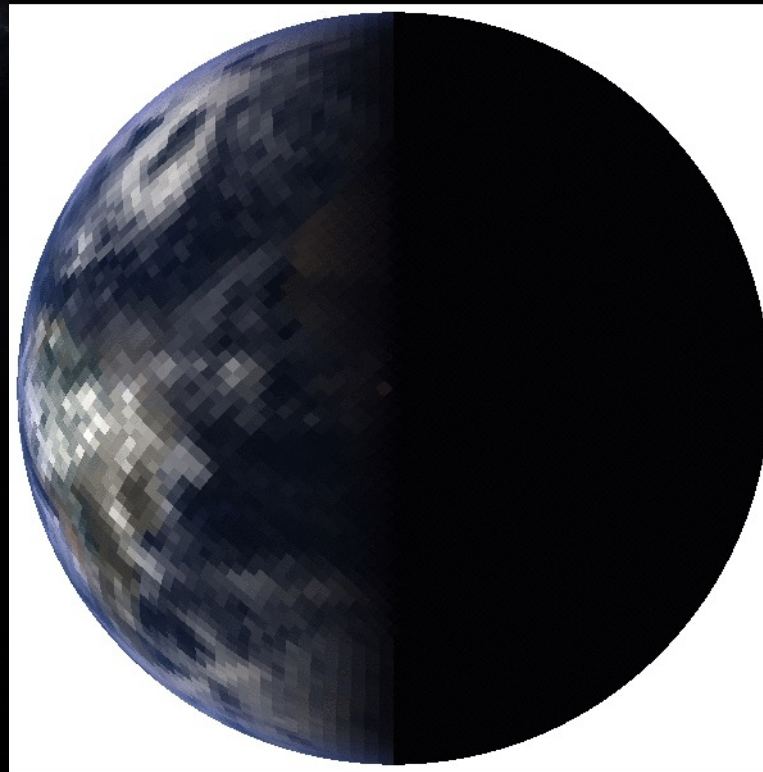
(Validated) Earth Models Can Provide Predictions at Other Phase Angles

Earth Orbital Phase Evolution

Gibbous ($\alpha = 45^\circ$)



Quadrature ($\alpha = 90^\circ$)



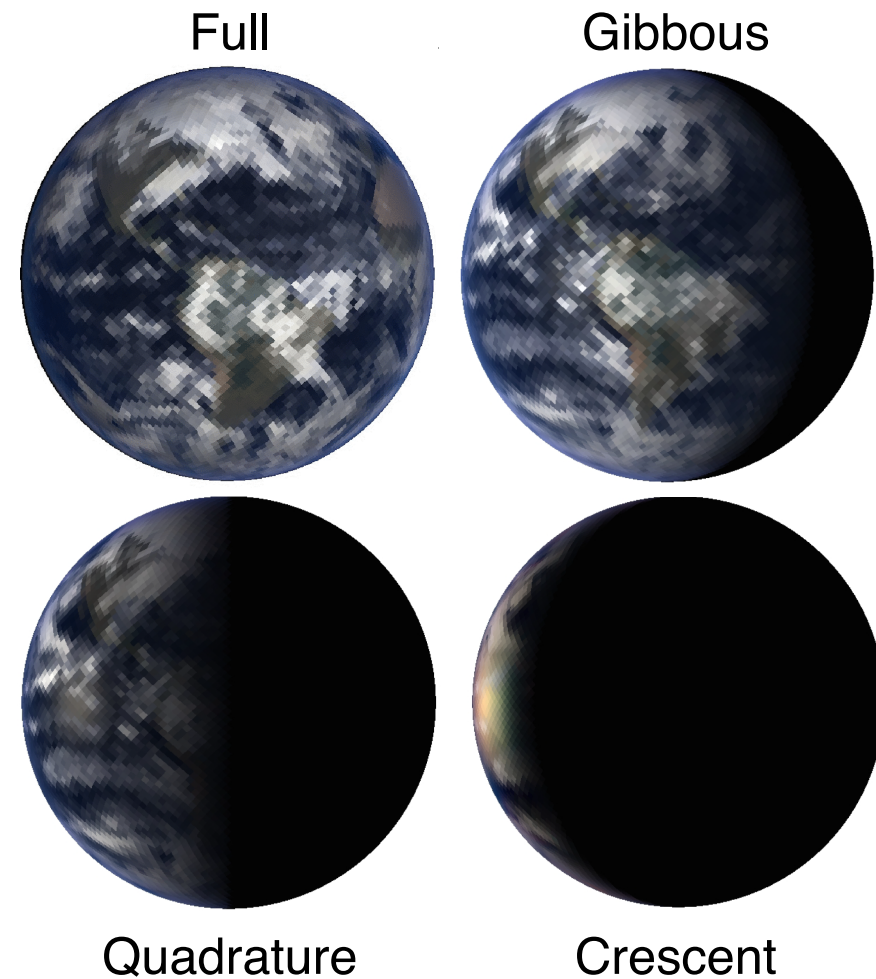
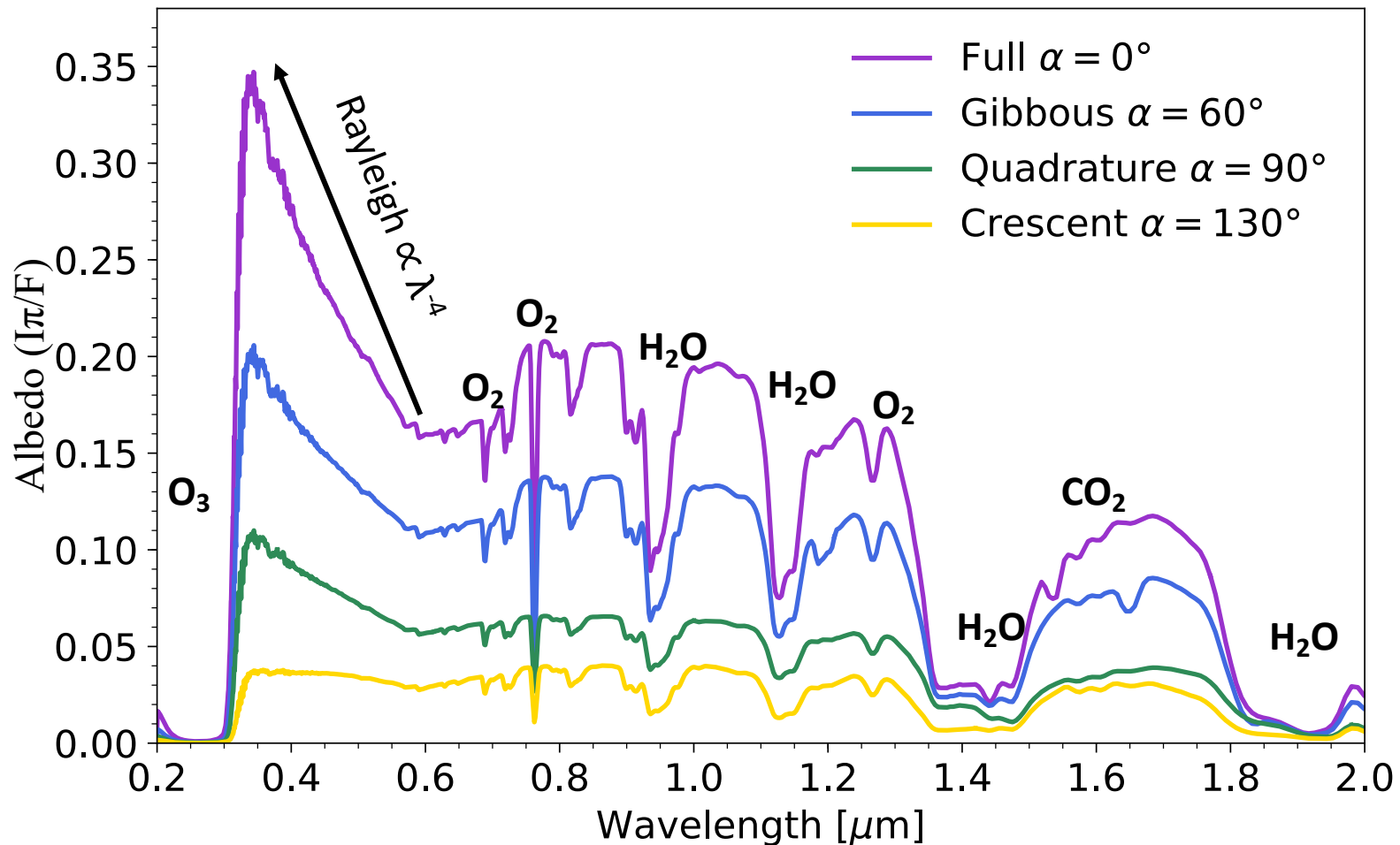
Crescent ($\alpha = 135^\circ$)



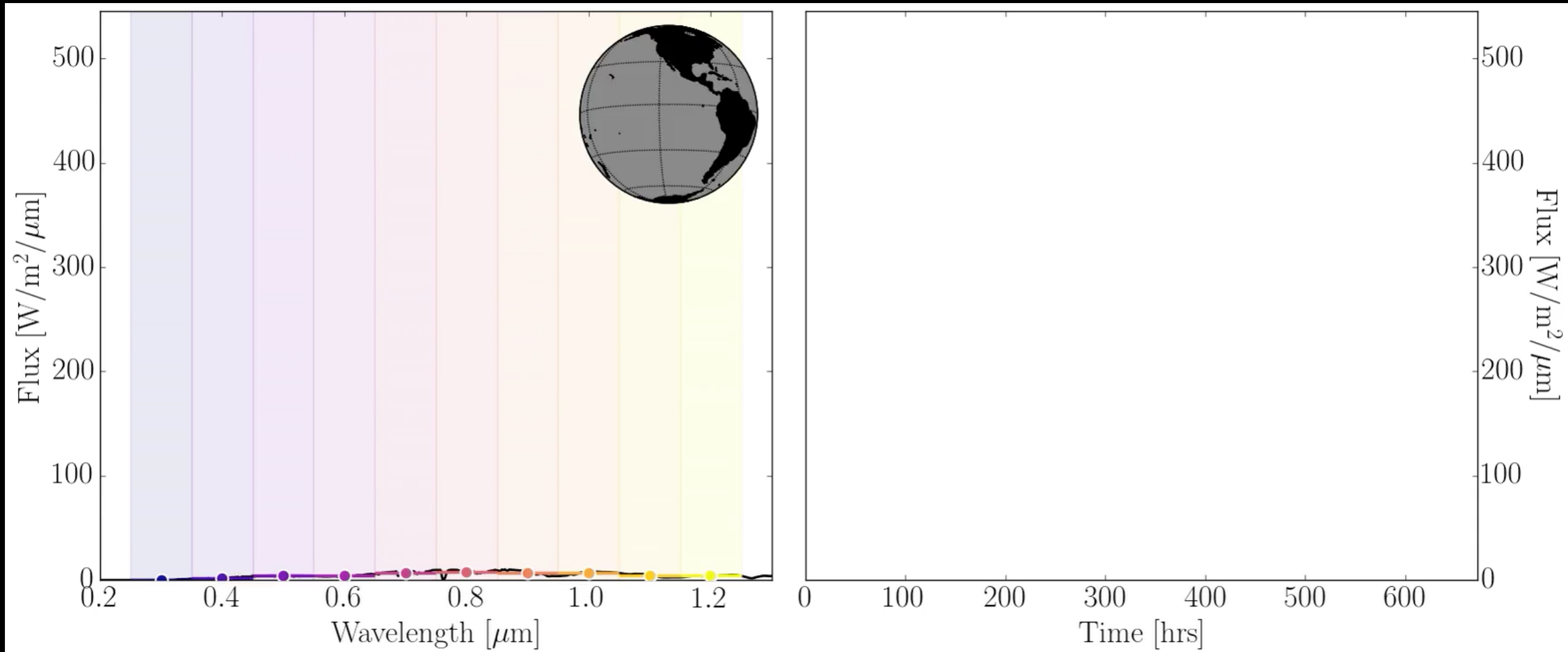
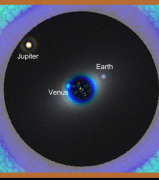
VPL Earth Model (e.g., Robinson+2011, 2014)

Spectrum as a Function of Phase

VPL Earth Model



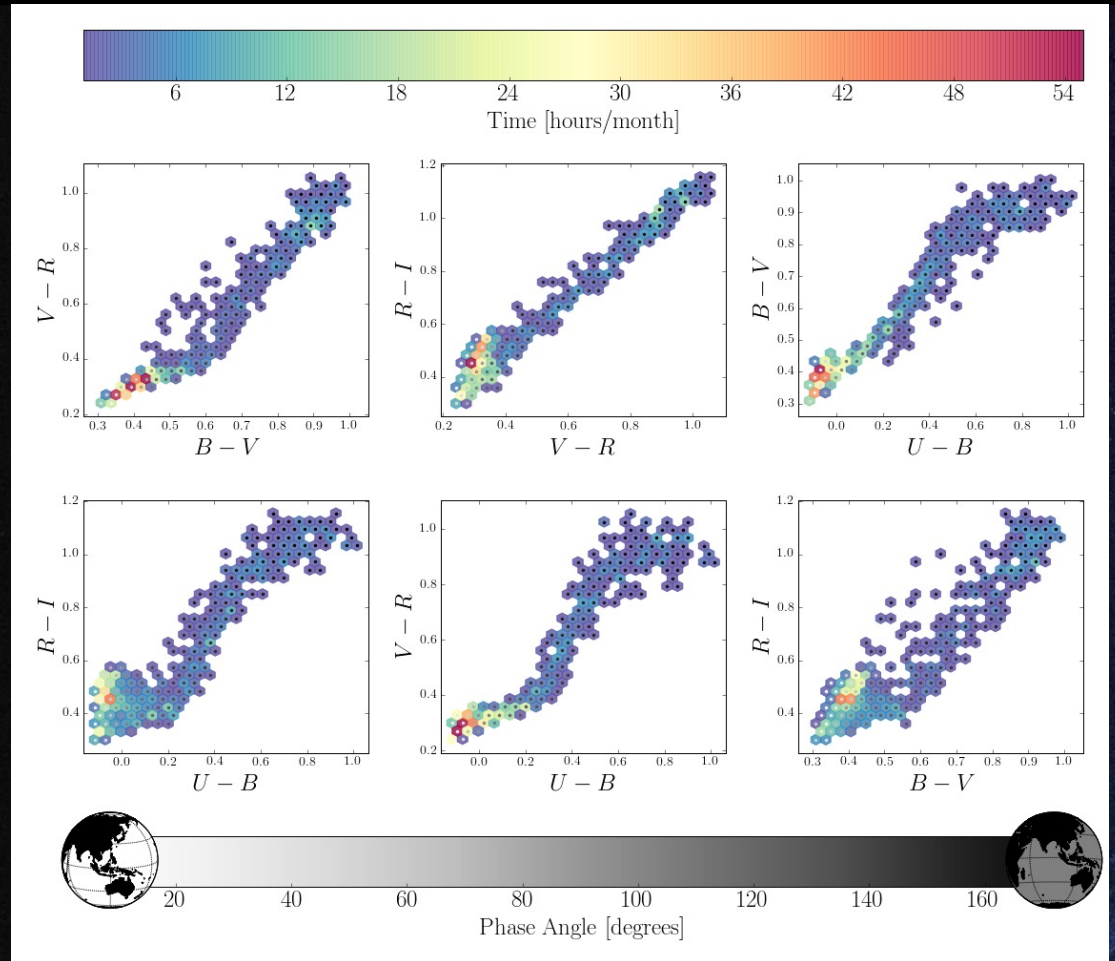
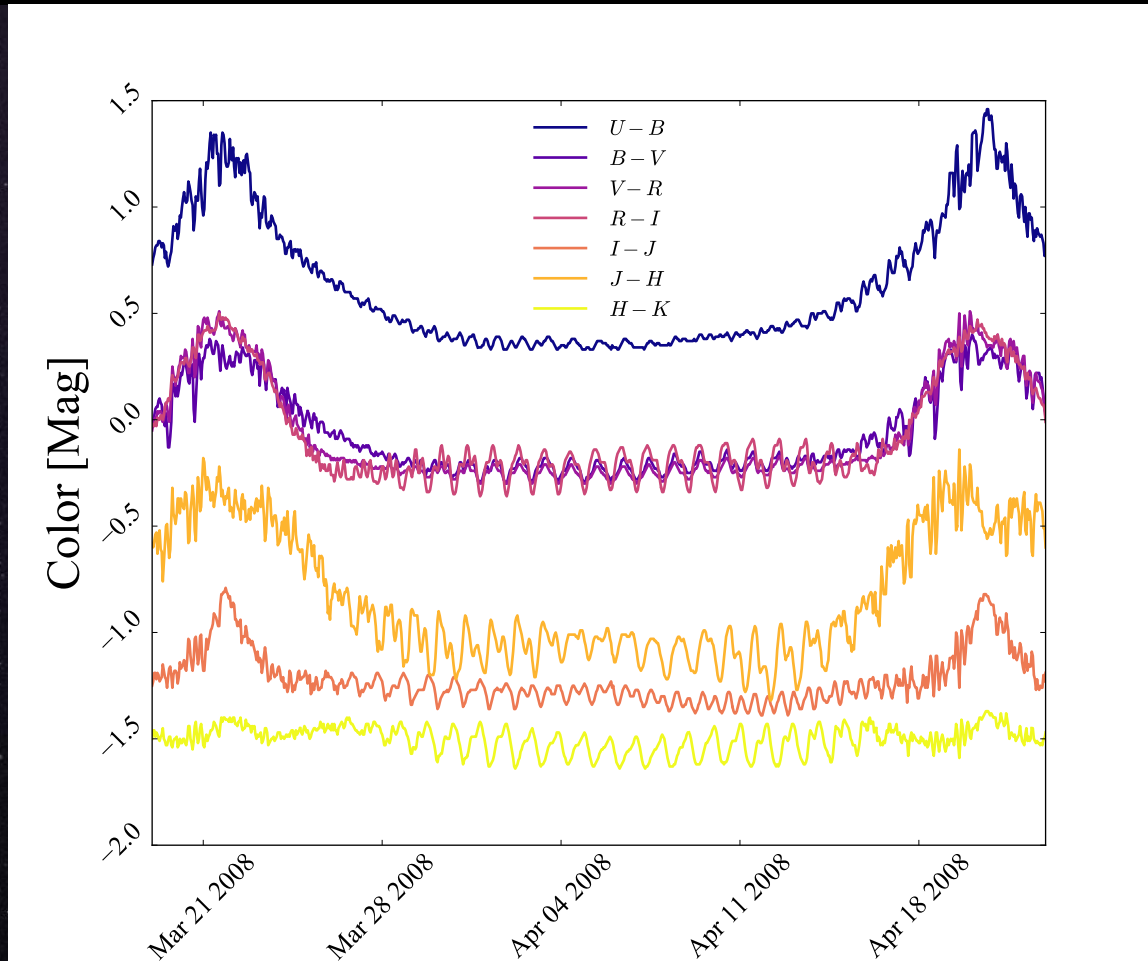
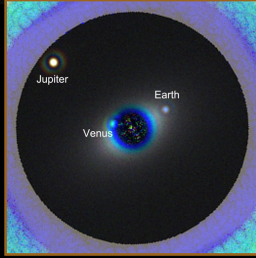
Earth Shows Wavelength-Dependent Phase and Diurnal Variations



Earth from Moon one month database with 1-hr cadence [Model]

Credit: J. Lustig-Yaeger, E. Schwieterman, & VPL Team

Earth Displays Diverse Colors Over Its Phase Evolution

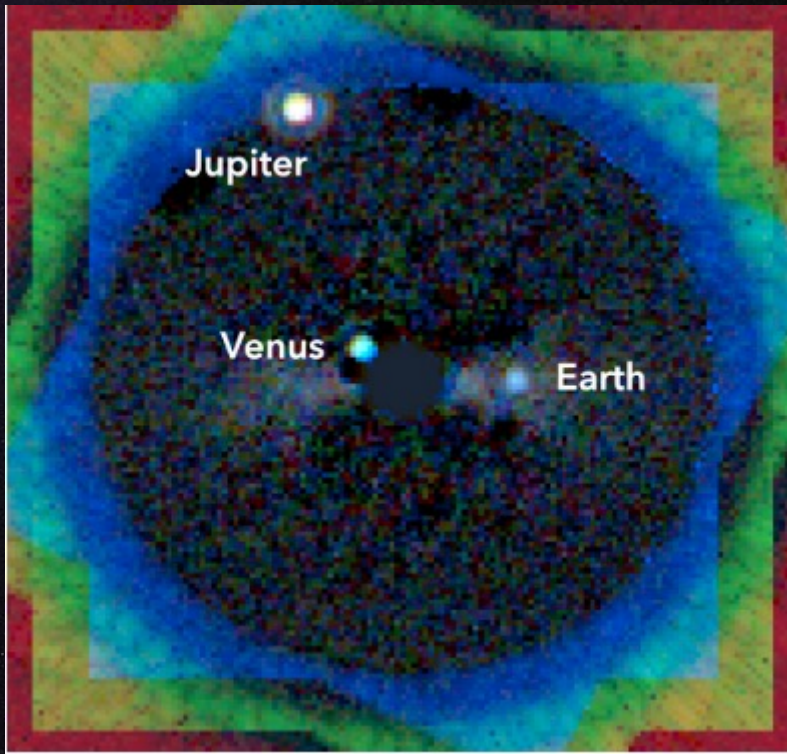


Earth from Moon one month database with 1-hr cadence [Model]

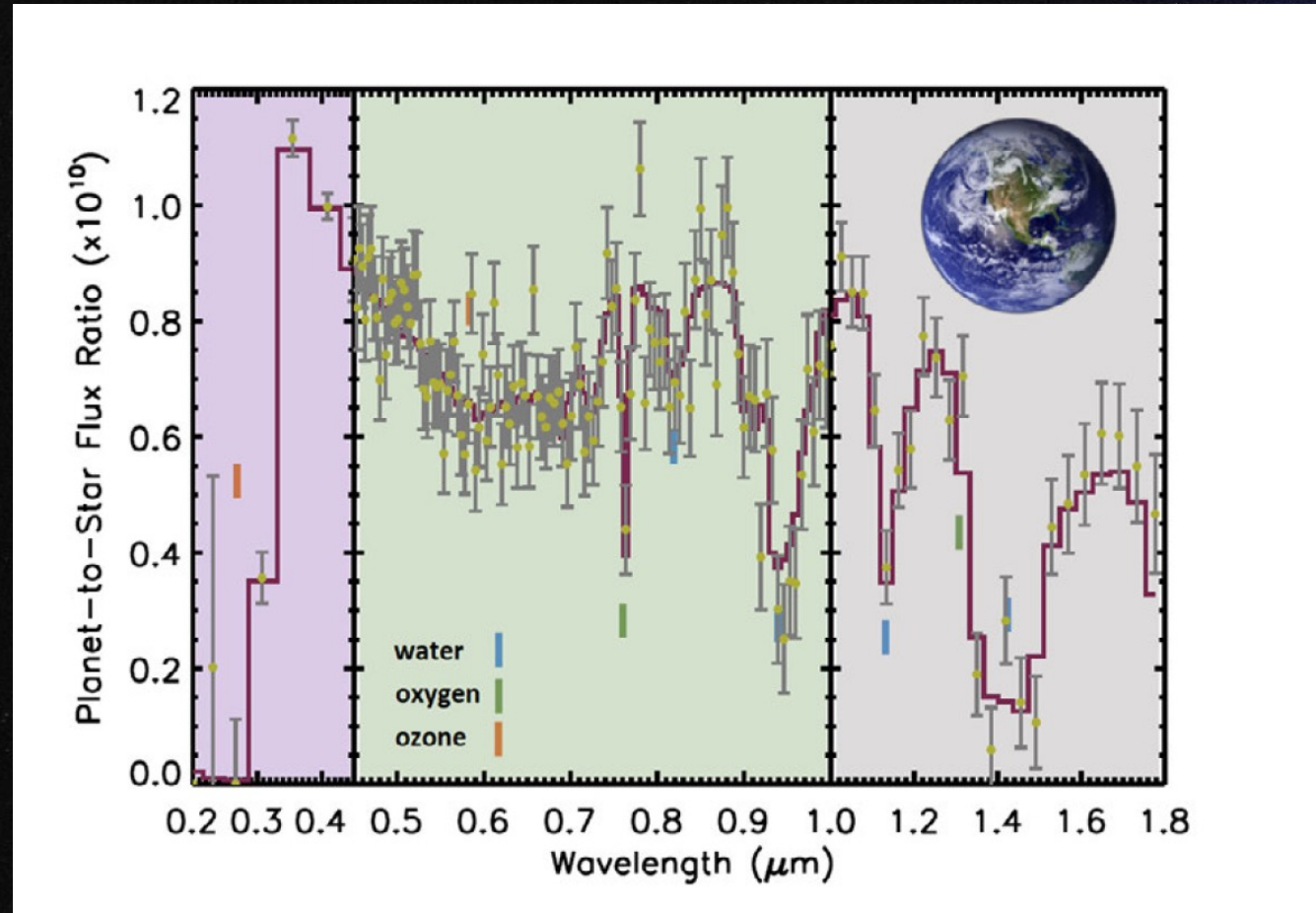
Credit: J. Lustig-Yaeger, E. Schwieterman, & VPL Team

How An ExoEarth Would Be Imaged in Reflected Light (HWO)

Star nulling by coronagraph or starshade (10^{10})

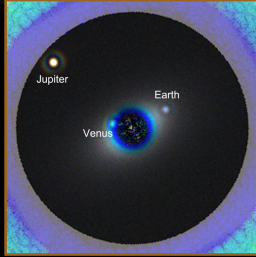


LUVOIR Report; Roberge et al. 2018

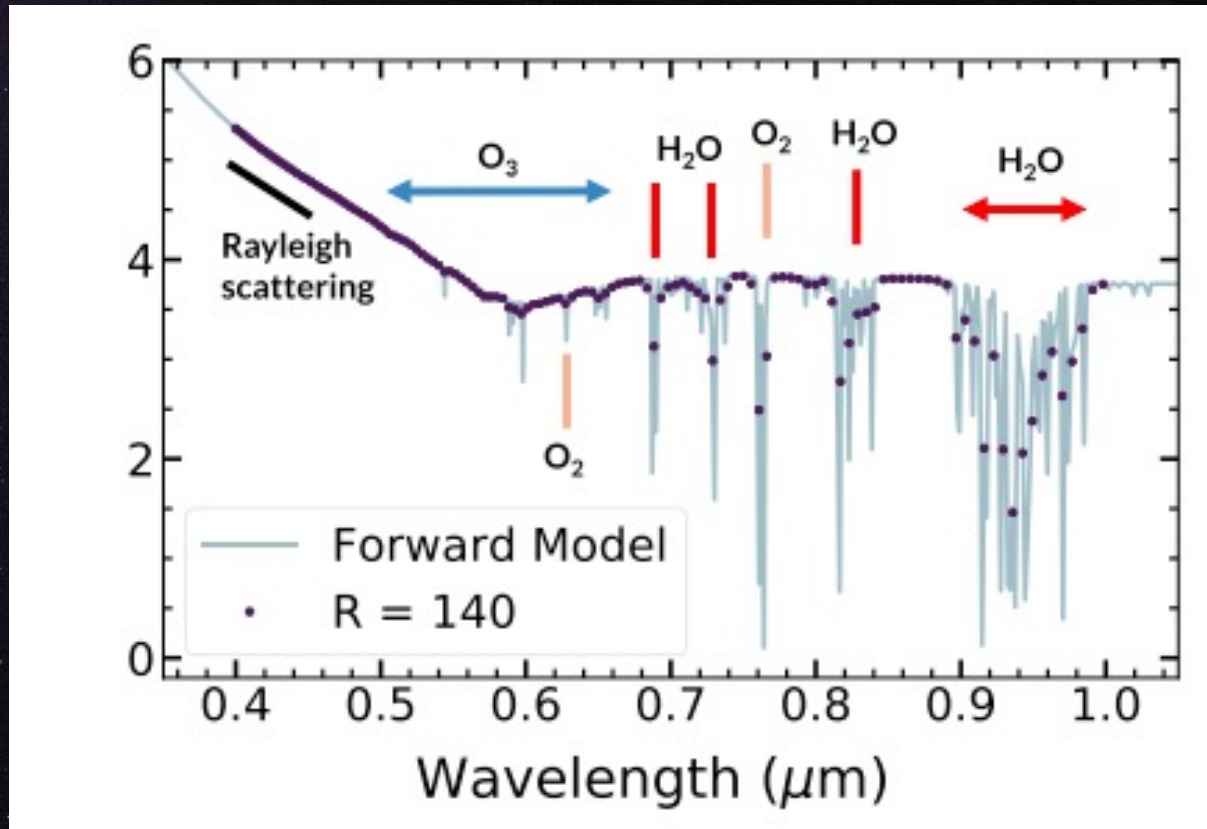


HabEx Report; Robinson et al. 2016

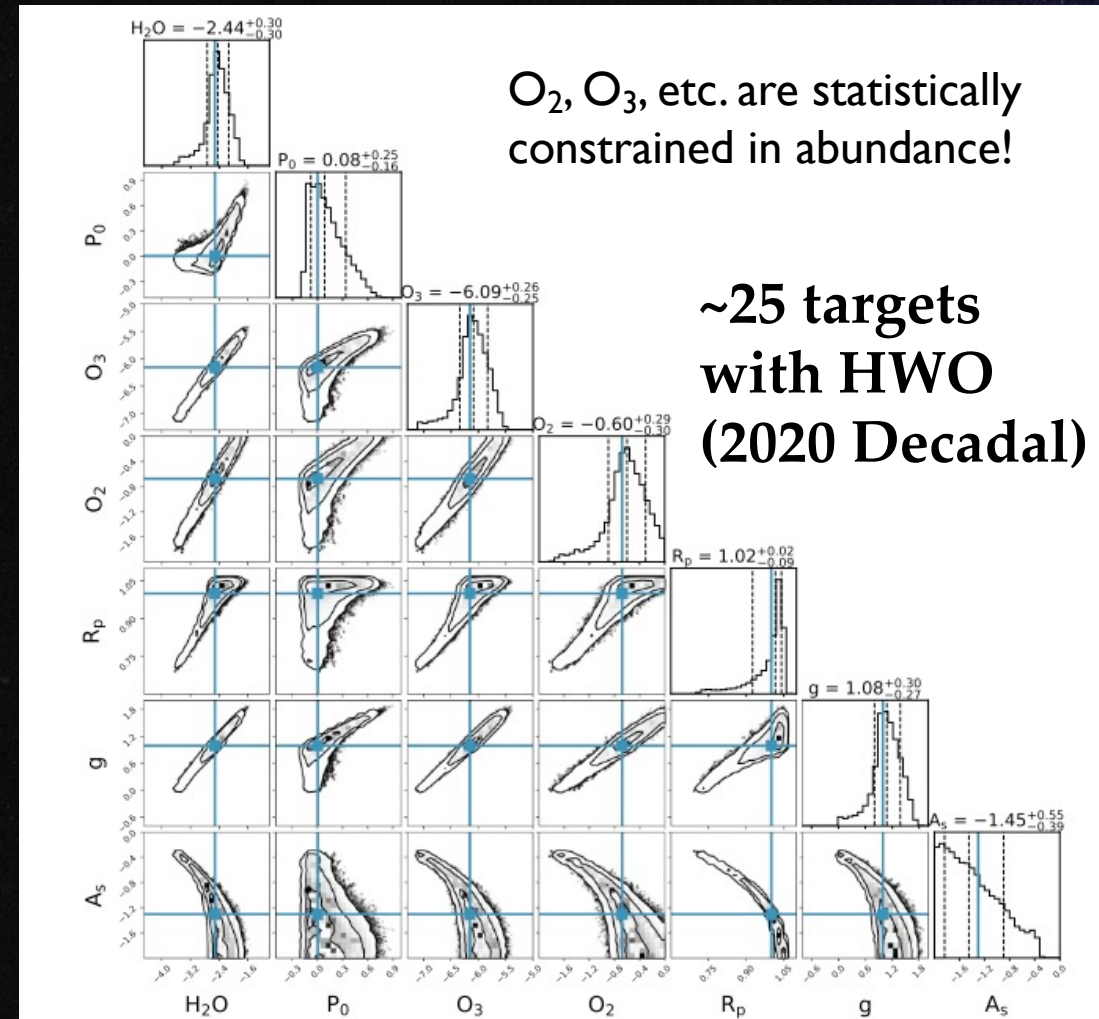
Atmospheric Retrievals Will Help Confirm Signatures



Employing an Inverse Model

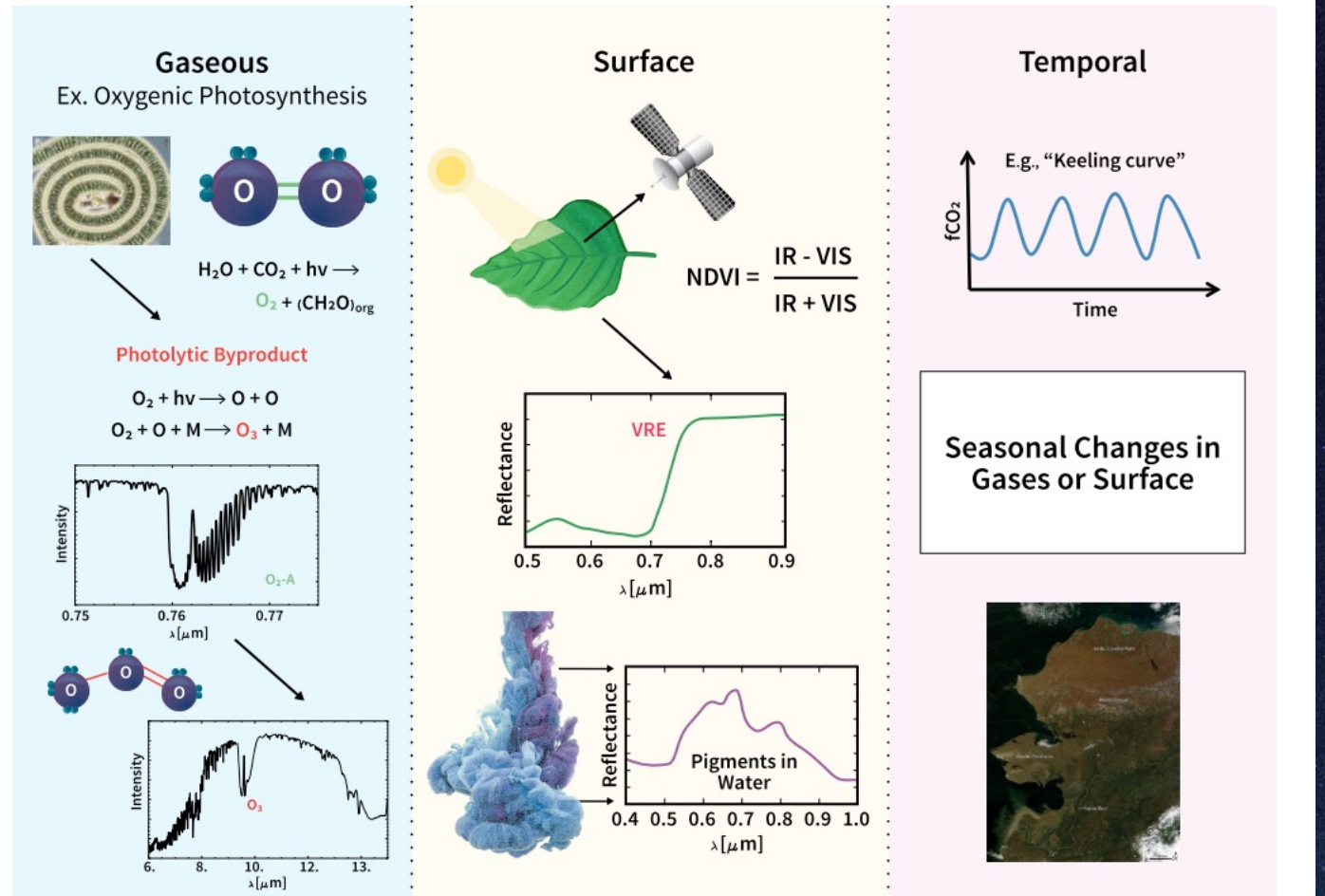


Feng et al. 2018, *AJ*, 155:200



Extending Earth Biosignatures to Those Possible on ExoEarths

- **Atmospheric Gases:** O₂, O₃, CH₄, N₂O, DMS—(CH₃)₂S, DMSe—(CH₃)₂Se, C₅H₈, CH₃Cl, CH₃Br, CH₃SH, PH₃, NH₃, etc.
- **Surface Features:** Vegetation Red Edge (VRE), anoxygenic photosynthesis, rhodopsins, other pigments
- **Temporal Changes:** Seasonal Change in Gas (e.g., CO₂, CH₄, O₃ or Pigments)
- **Always Context-Dependent!**



Schwieterman et al., 2018, 2021

Earth Through Time: Biosignature Combinations (Reflected Light)

Anoxic, no haze: $\text{CH}_4 + \text{H}_2\text{O} + \lambda^{-4} + \text{CO}_2$

Strongest Anoxic: haze + $\text{H}_2\text{O} + \text{CH}_4 + \text{CO}_2$

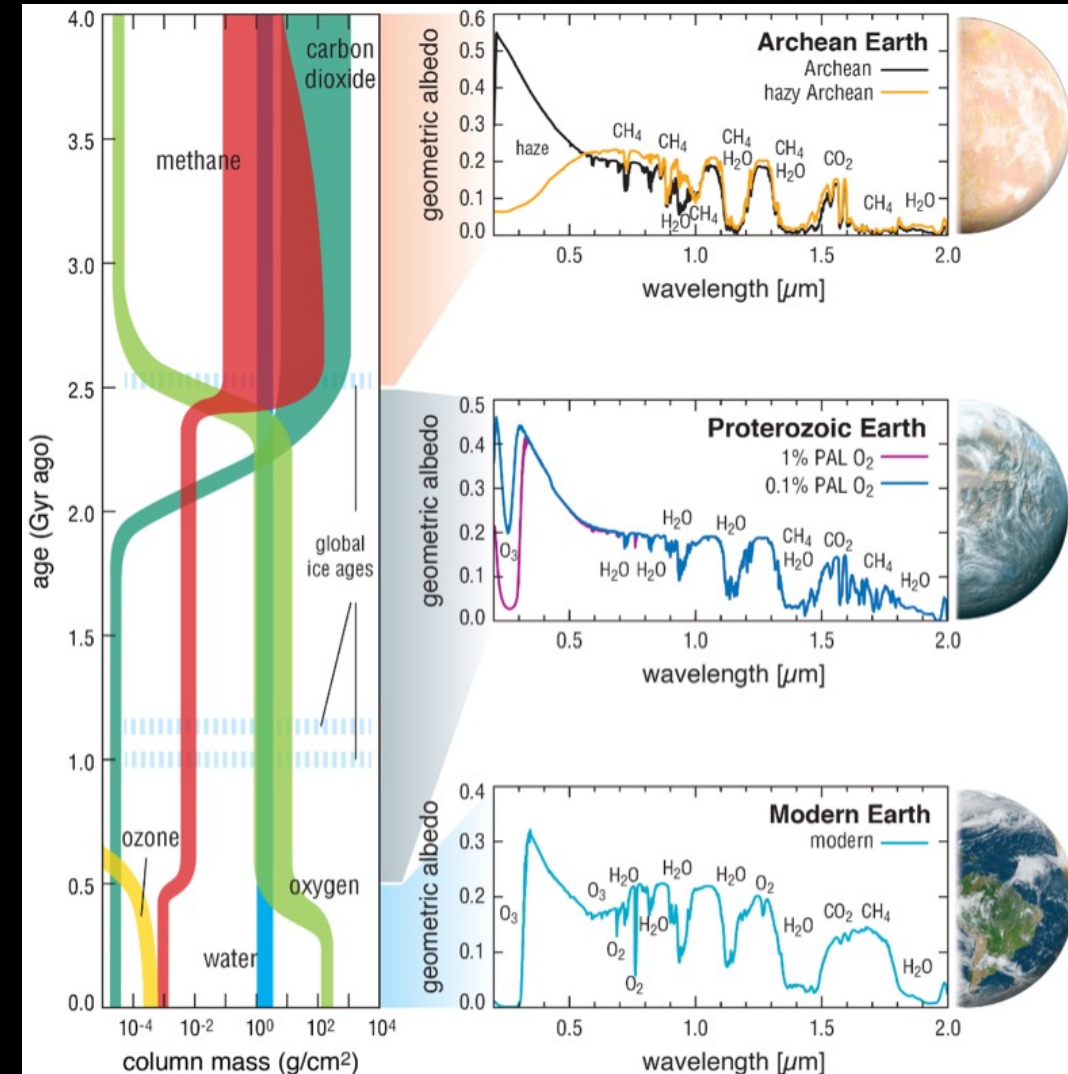
Weakly Oxygenated: $\text{O}_3 + \lambda^{-4} + \text{H}_2\text{O}$ [+ CH_4]

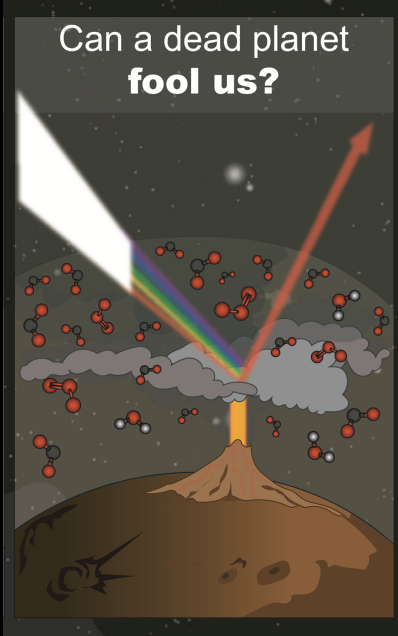
Strong Oxygen-rich: $\text{O}_3 + \lambda^{-4} + \text{H}_2\text{O} + \text{O}_2$

Strongest Oxygen-rich: $\text{O}_3 + \lambda^{-4} + \text{H}_2\text{O} + \text{O}_2 + \text{CH}_4$

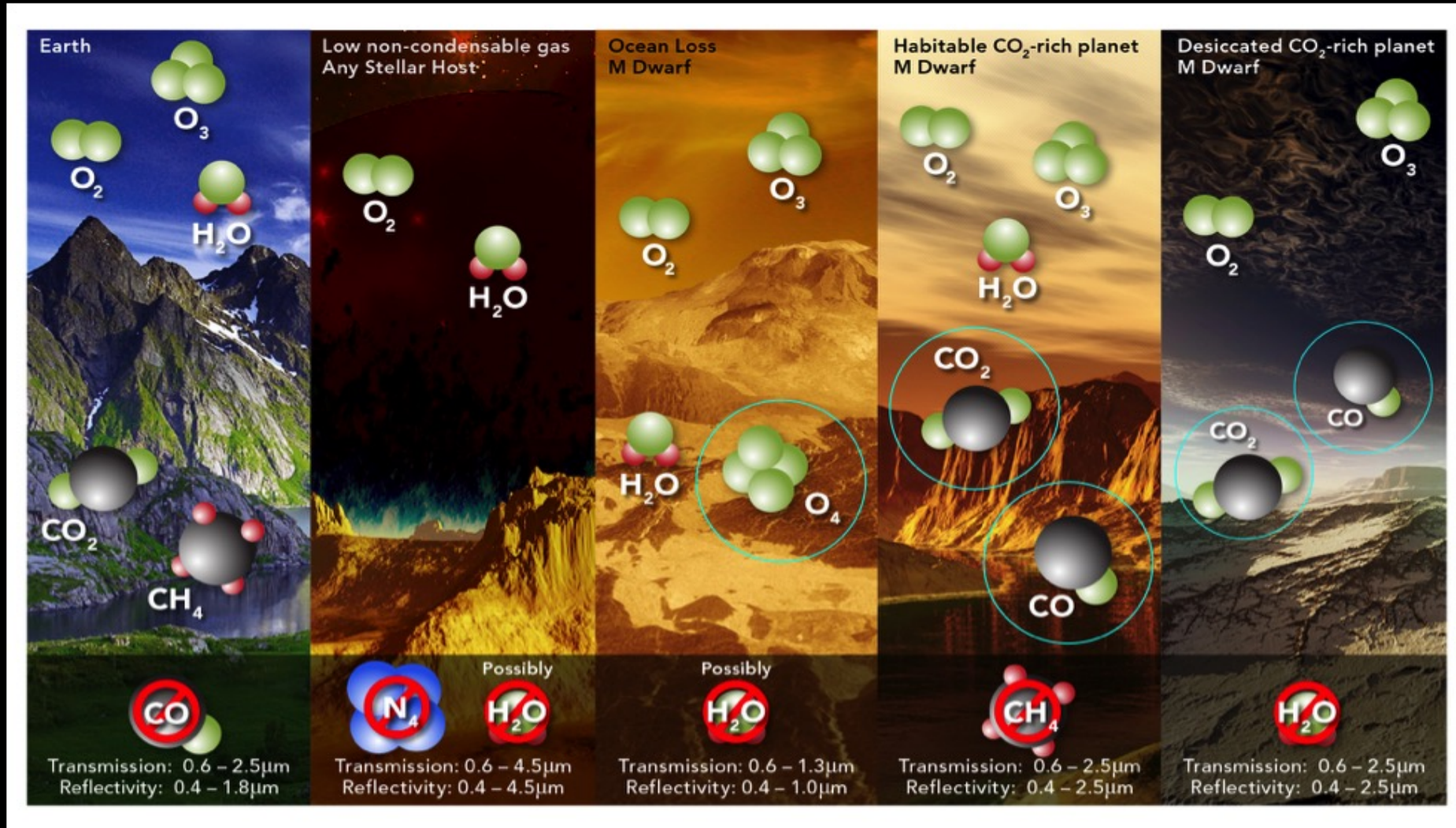
Biosignatures and false positive evaluators are ordered by their relative detectability in atmospheric spectra

LUVOIR Final Report





Context can help rule out “false positives”



Detecting H₂O, CH₄, and/or the Rayleigh slope ($\propto \lambda^{-4}$) can confirm biological nature of O₂/O₃.

Certain features (O₂-O₂ CIA) would confirm a “false positive.”

Image Credit R. Hasler e.g., Tian+2015; Domagal-Goldman+14, Hu+2020; Meadows+2017,2018; Schwieterman+2016

Searching for Life via Thermodynamic Disequilibria

A search for life on Earth from the Galileo spacecraft

Carl Sagan*, W. Reid Thompson*, Robert Carlson†, Donald Gurnett‡ & Charles Hord§

TABLE 1 Constituents of the Earth's atmosphere (volume mixing ratios)

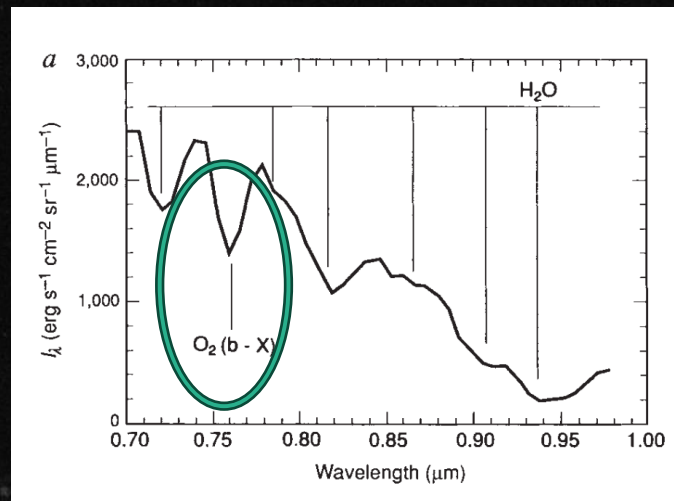
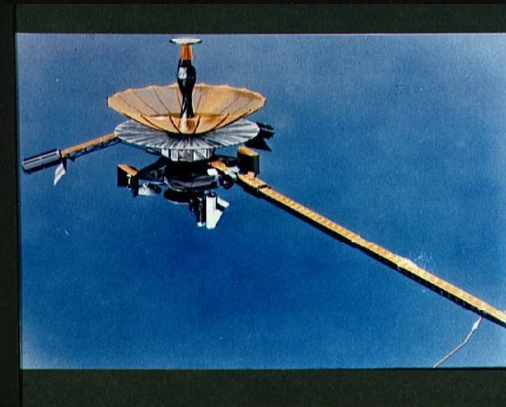
Molecule	Standard abundance (ground-truth Earth)	Galileo value*	Thermodynamic equilibrium value	
			Estimate 1†	Estimate 2‡
N ₂	0.78		0.78	
O ₂	0.21	0.19 ± 0.05	0.21§	
H ₂ O	0.03–0.001	0.01–0.001	0.03–0.001	
Ar	9 × 10 ⁻³		9 × 10 ⁻³	
CO ₂	3.5 × 10 ⁻⁴	5 ± 2.5 × 10 ⁻⁴	3.5 × 10 ⁻⁴	
CH ₄	1.6 × 10 ⁻⁶	3 ± 1.5 × 10 ⁻⁶	< 10 ⁻³⁵	10 ⁻¹⁴⁵
N ₂ O	3 × 10 ⁻⁷	~10 ⁻⁶	2 × 10 ⁻²⁰	2 × 10 ⁻¹⁹
O ₃	10 ⁻⁷ –10 ⁻⁸	> 10 ⁻⁸	6 × 10 ⁻³²	3 × 10 ⁻³⁰

* Galileo values for O₂, CH₄ and N₂O from NIMS data; O₃ estimate from UVS data.

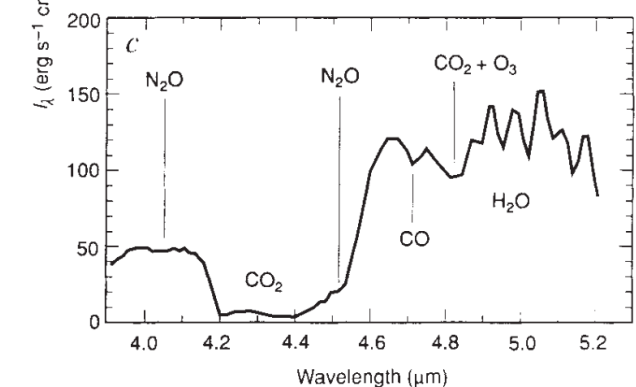
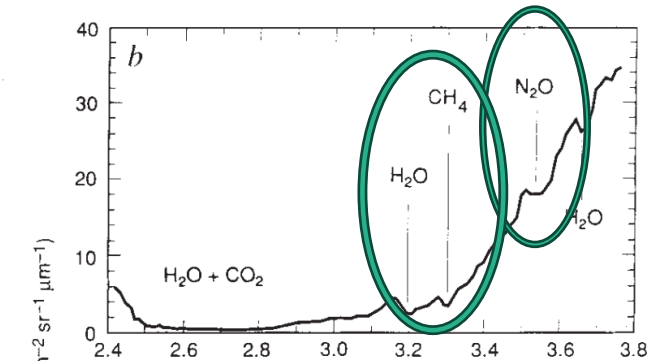
† From ref. 16 (P, 1 bar; T, 280 K).

‡ From ref. 17 (P, 1 bar; T, 298 K).

§ The observed value; it is in thermodynamic equilibrium only if the under-oxidized state of the Earth's crust is neglected.



Kinetic disequilibrium is most important! CH₄ lifetime ~10 years. N₂O ~140 years.



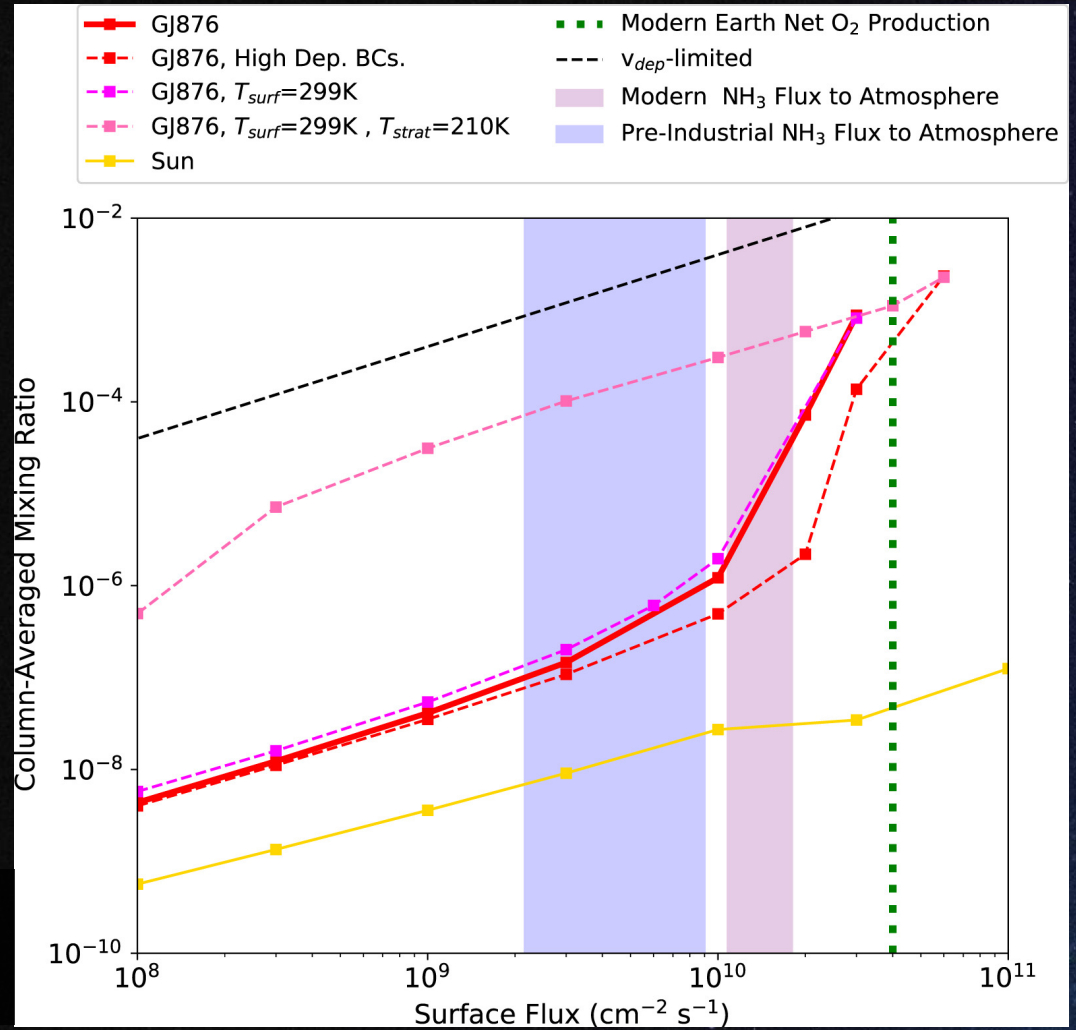
Biogenic Gases on Earth: “Equilibrium” vs. “Non-Equilibrium” Gases

“Photochemical Runaway” (Ranjan+2022, AJ, 930:131)

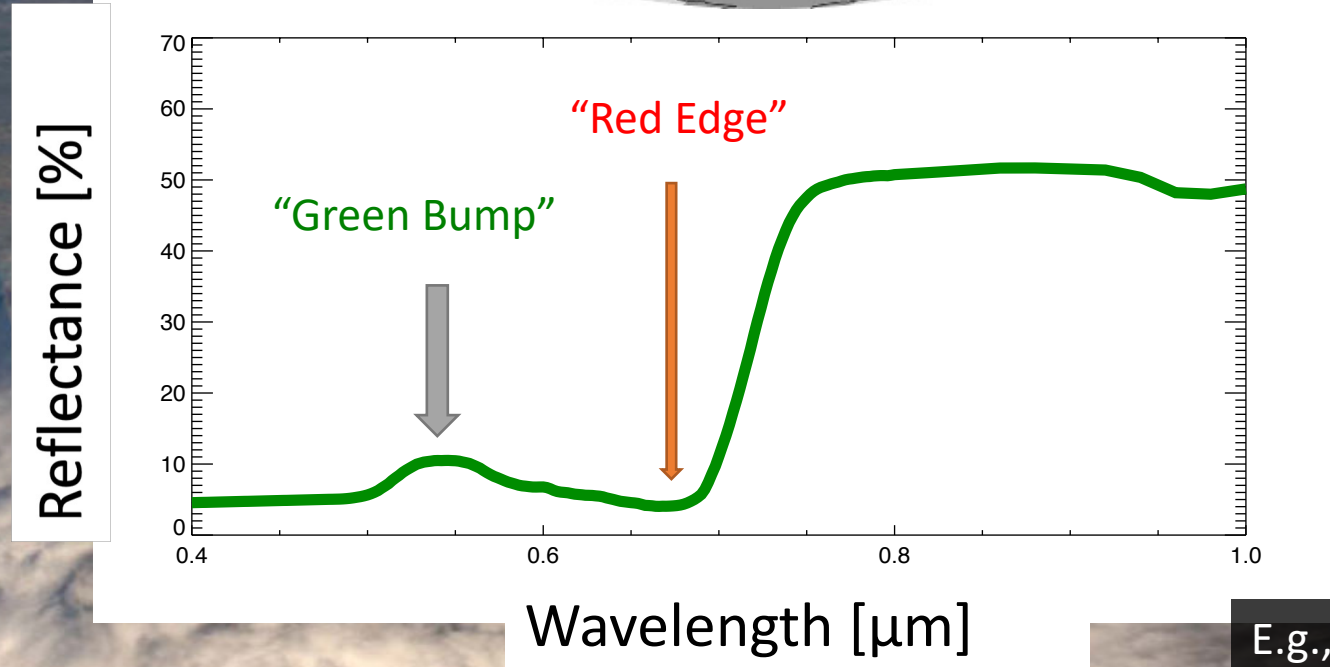
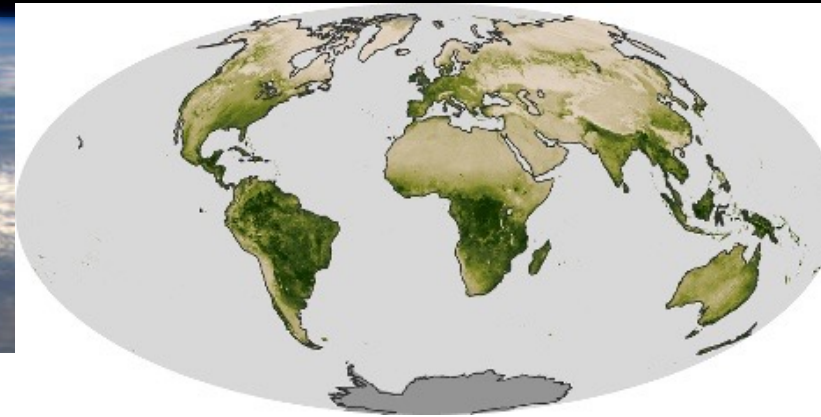
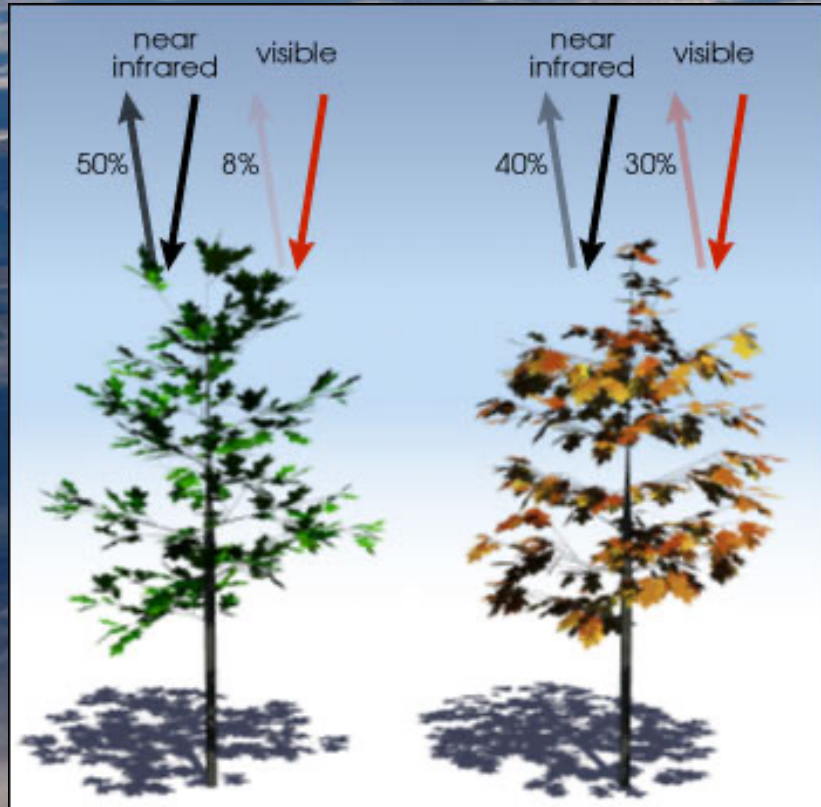
Gases formed at equilibrium in gas giant atmospheres (or via volcanism), but with strong kinetic disequilibrium in N_2 - CO_2 - $[-O_2]$ or temperate H_2 Atmospheres, e.g., CH_4 (Thompson+2022), NH_3 (Philips+2021, Huang+2022), PH_3 (Sousa-Silva+2020).

Gases not formed via equilibrium (albeit with other abiotic sources), e.g., $DMS/DMDS$ (Domagal-Goldman+2011), isoprene — C_5H_8 (Zhan+2021), $halomethanes$ (Leung+2022; see poster!), N_2O (Schwieterman+2022).

Gas lifetime will be a strong function of flux and host star spectrum — longer lifetimes for later stars.



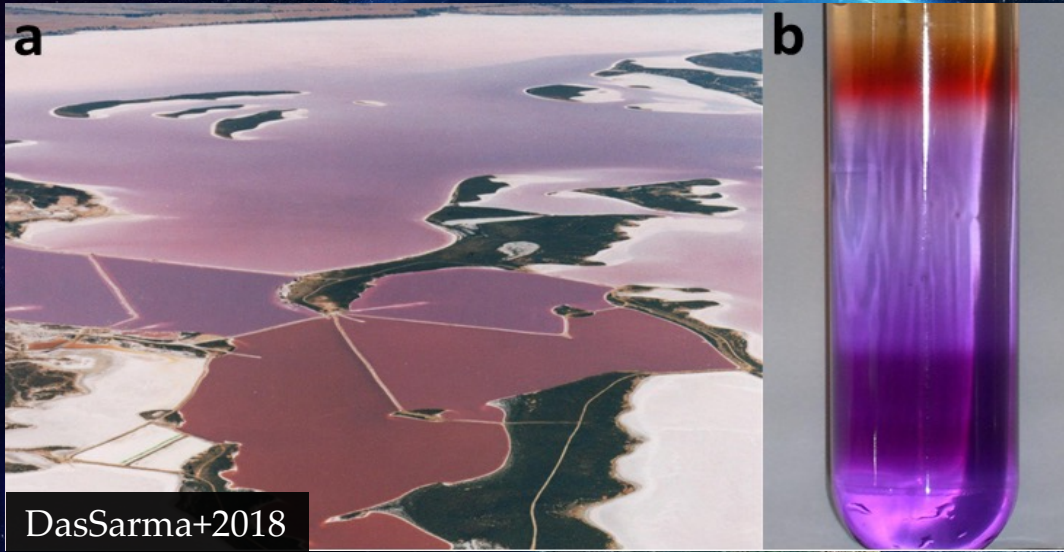
Surface Biosignature: Vegetation “Red Edge”



NASA Earth Observatory

E.g., Seager (2005)

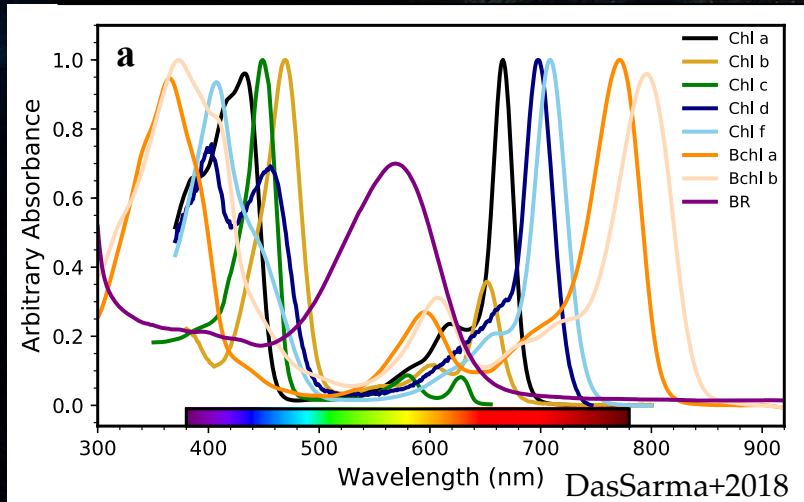
Halophiles: Rhodopsin & Carotenoids



Bacteriochlorophylls and Carotenoids



Complementary Absorption

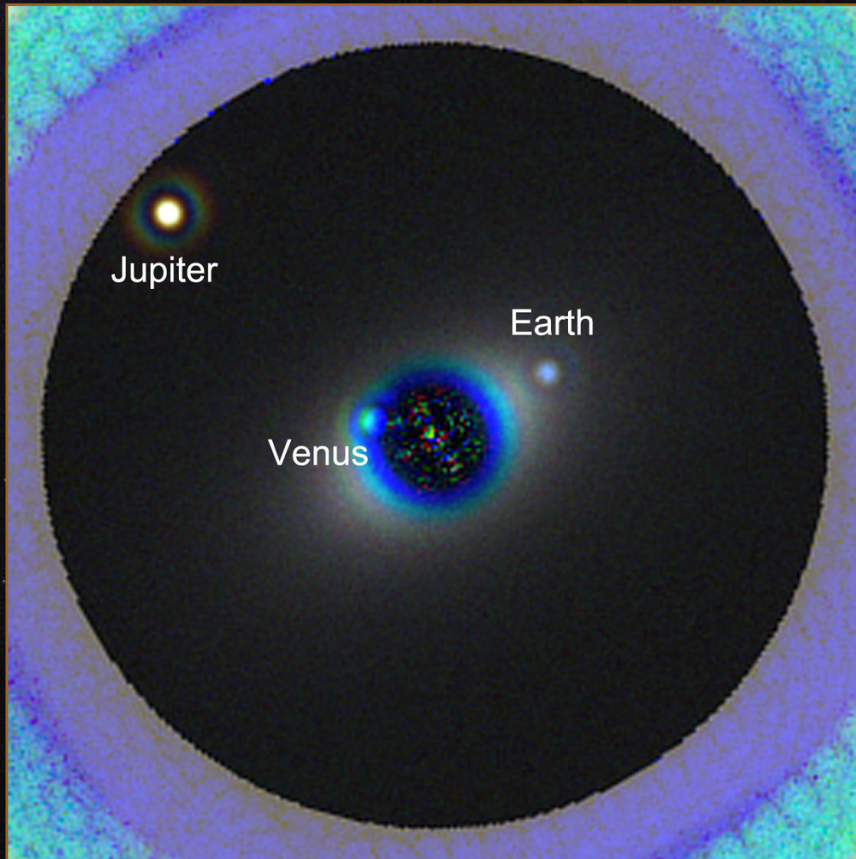


San Francisco Salt Ponds



Questions?

Thank you!



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- ICAR grants 80NSSC21K0594 & 80NSSC21K0905
- NASA Astrobiology Program grant 80NSSC18K082
- NASA XRP grant 80NSSC22K0235

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