

# Carbon Cycling on the TRAPPIST-1 Planets

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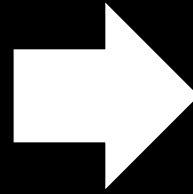
2023 NASA ExoExplorer Series

Johns Hopkins University

Advisor: Prof. Laura Schaefer (Stanford University)



# From Puerto Rico to Maryland





# Our Home: Planet Earth

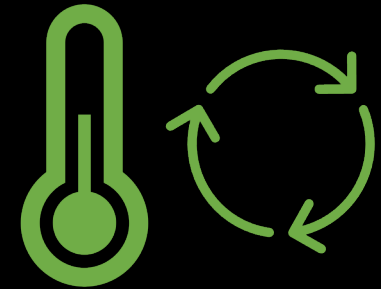


# Our Home: Planet Earth

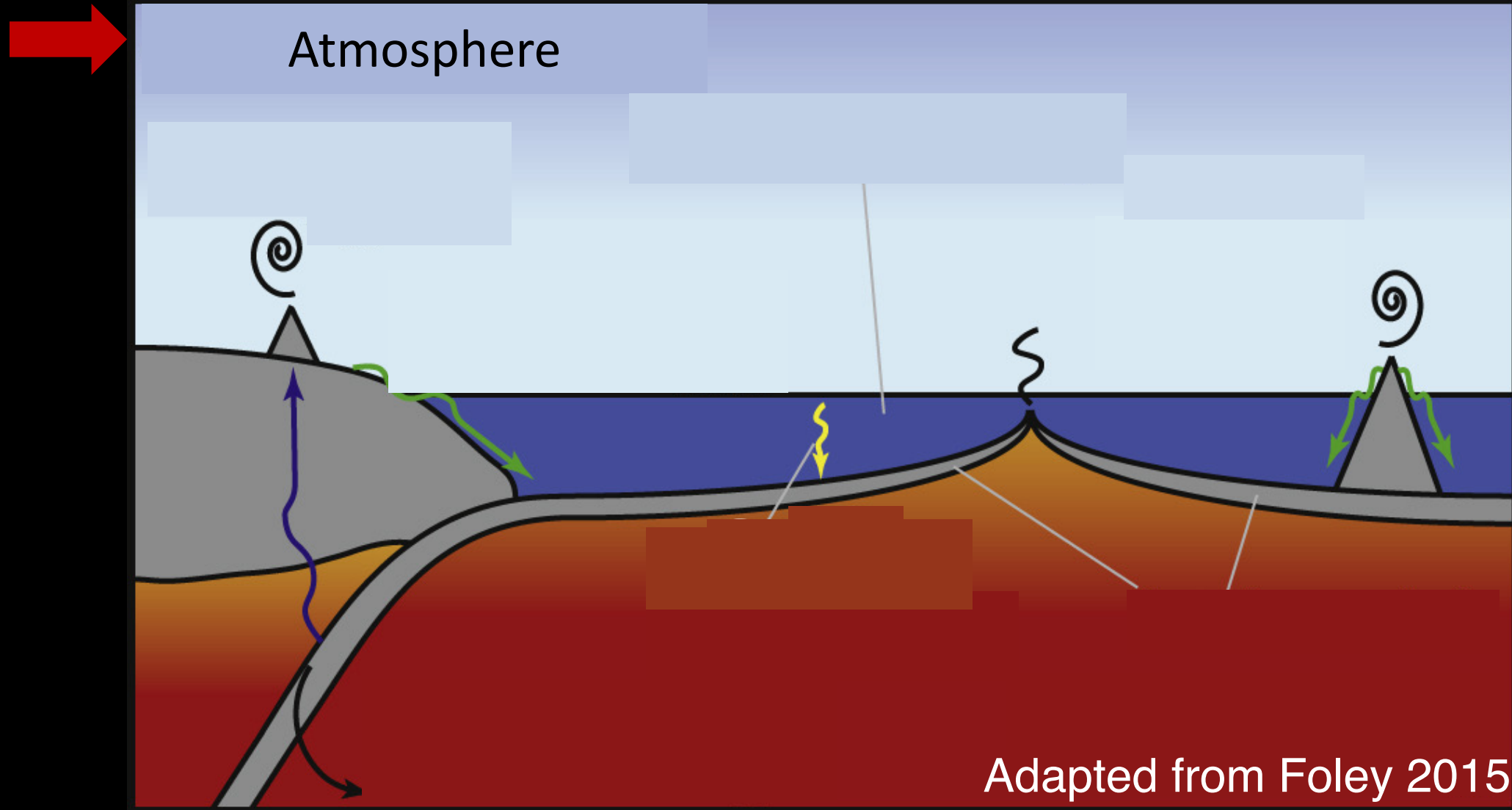
Has plate tectonics!



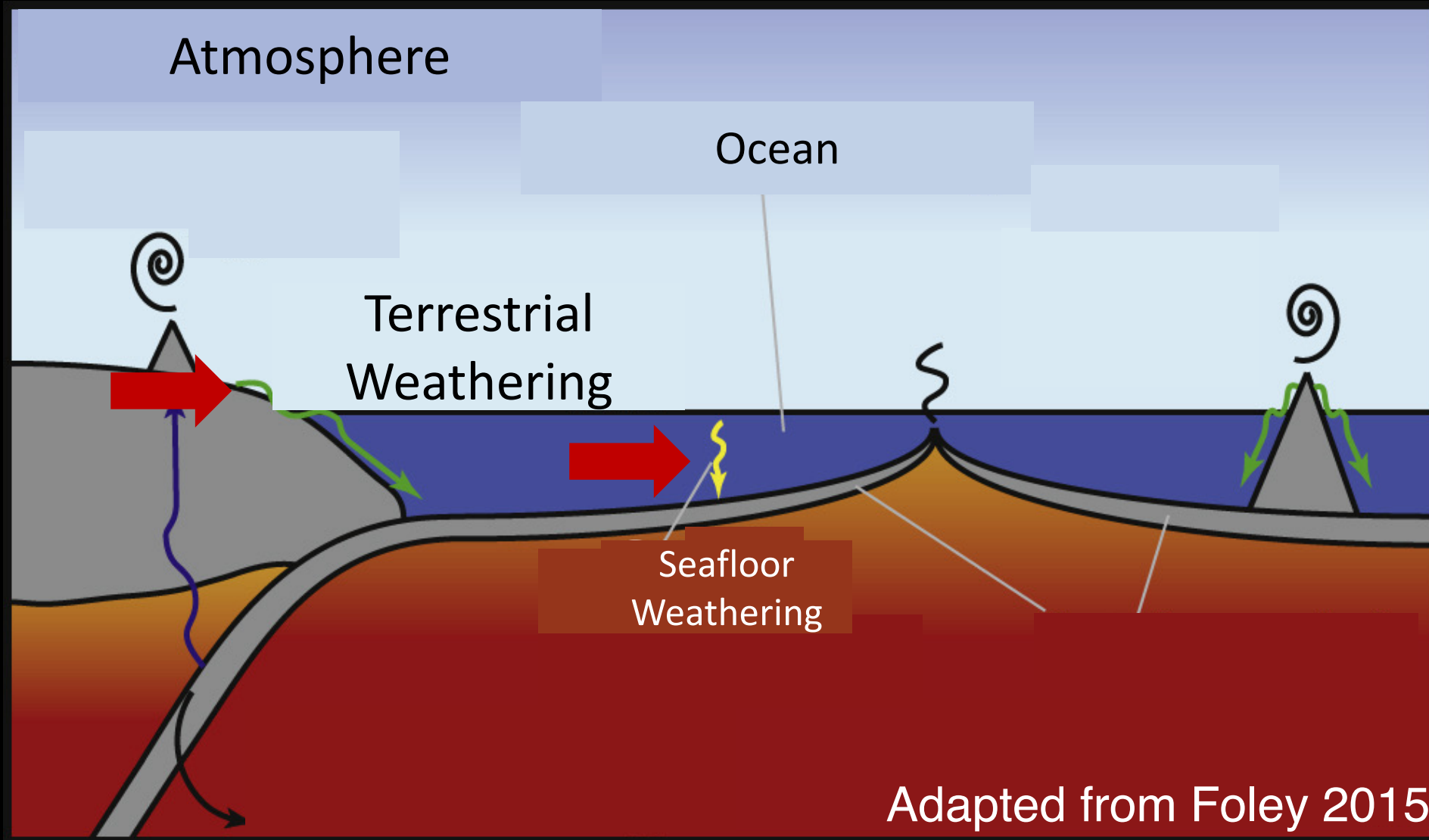
Drive important  
geological processes  
occurring such as the  
carbon cycle



# The carbon cycle regulates the climate of a planet over long timescales

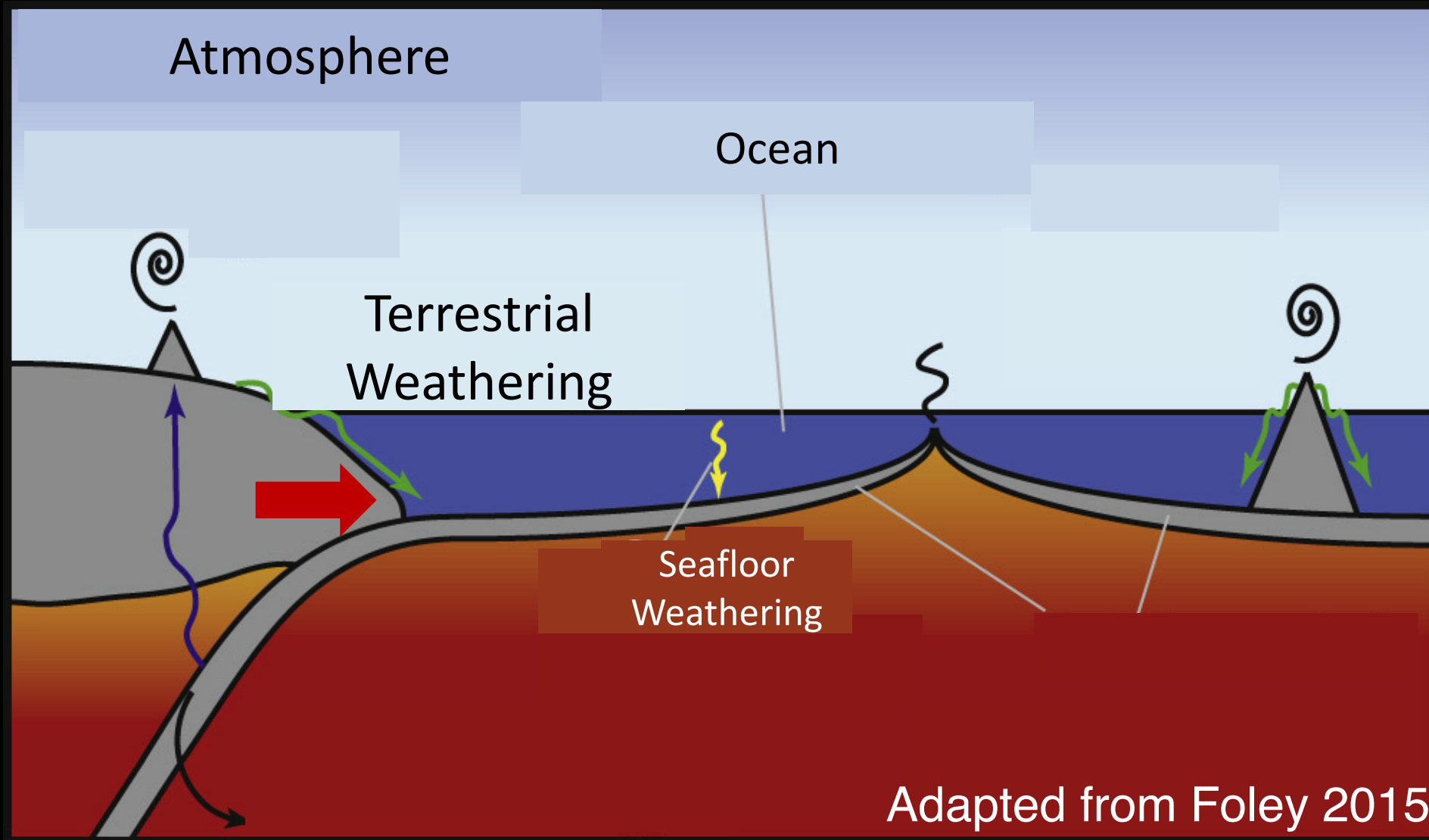


# The carbon cycle regulates the climate of a planet over long timescales

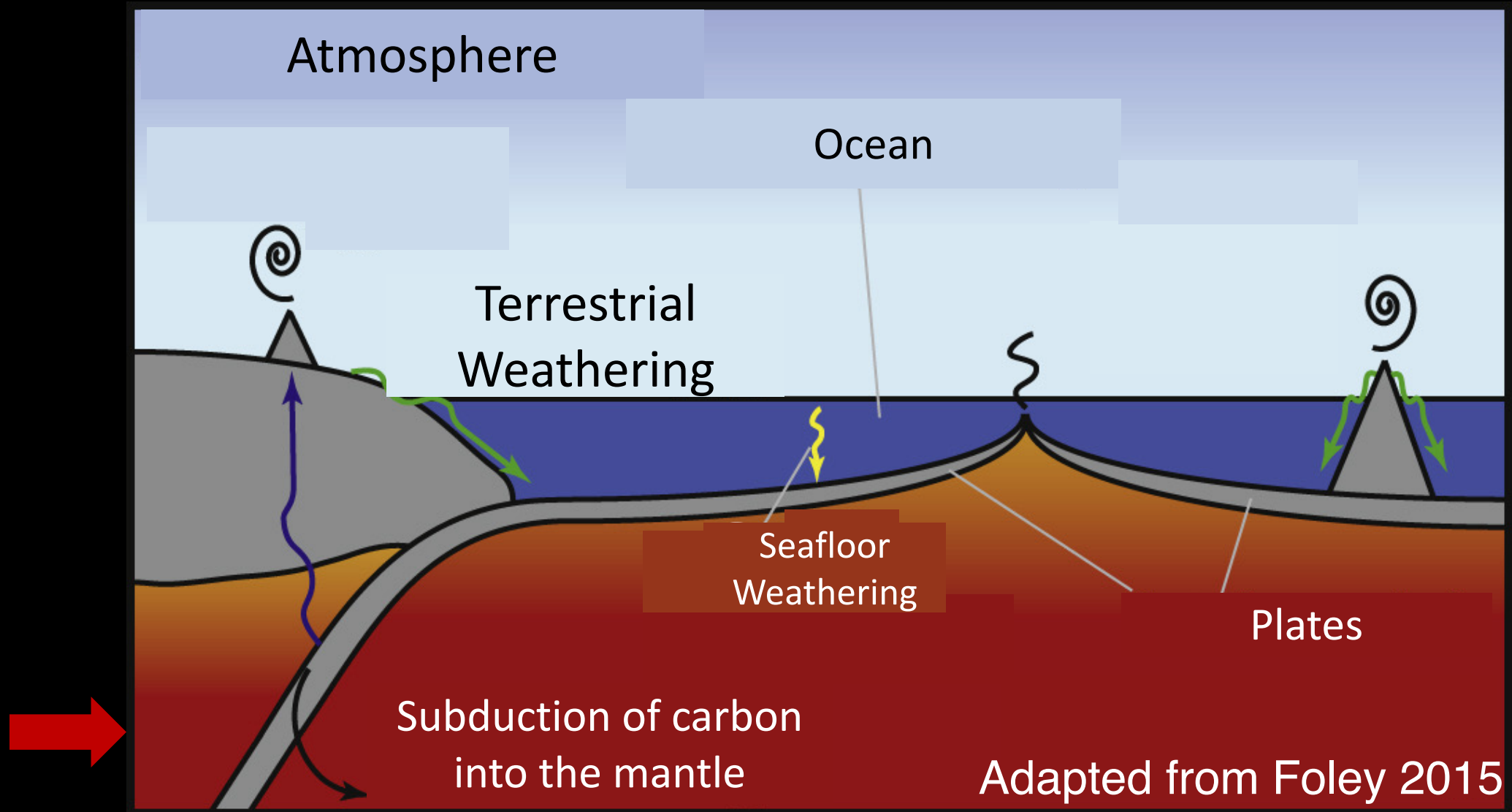




# The carbon cycle regulates the climate of a planet over long timescales

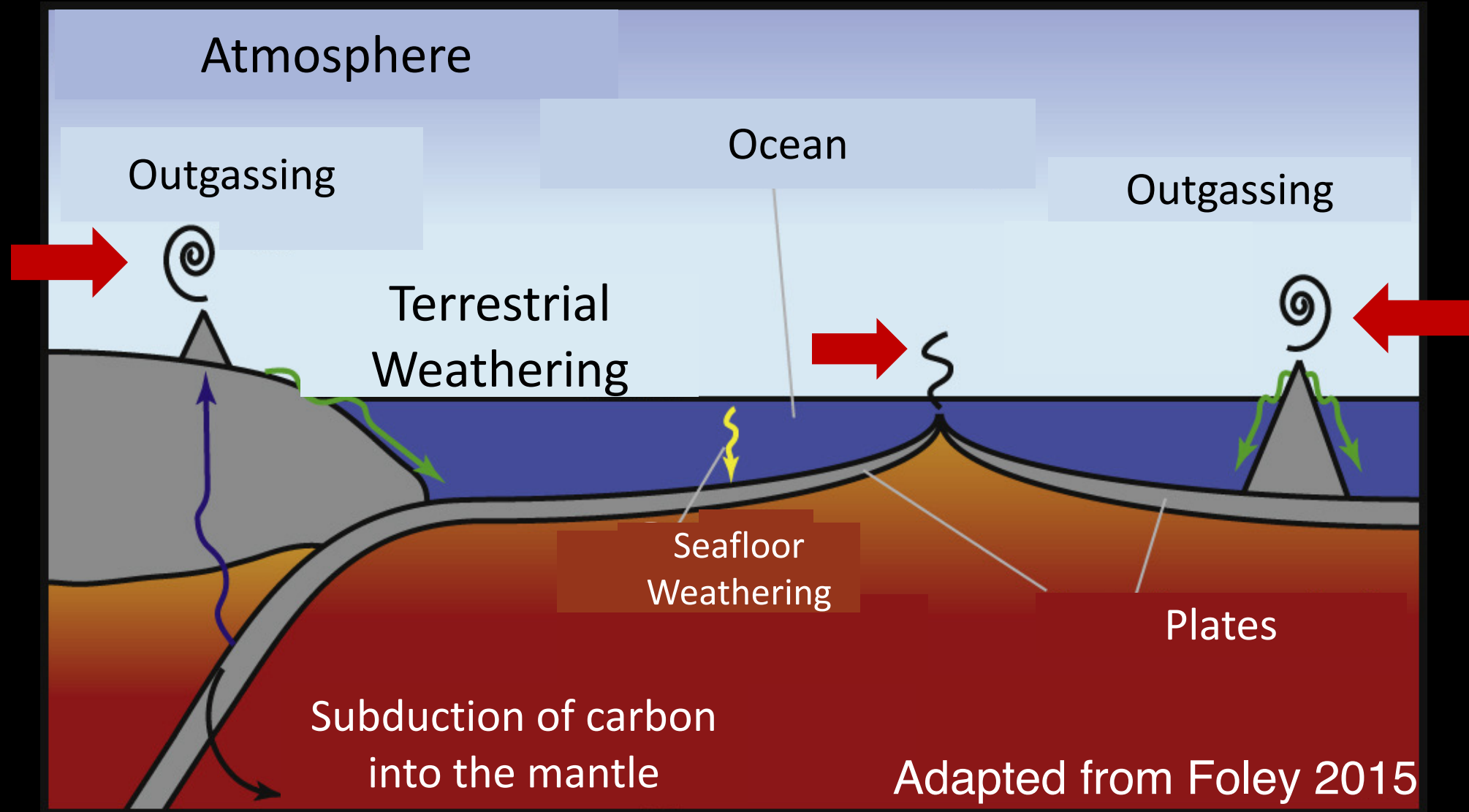


# The carbon cycle regulates the climate of a planet over long timescales





# The carbon cycle regulates the climate of a planet over long timescales



# The carbon cycle regulates the climate of a planet over long timescales

Atmosphere

Ocean

**Regassing or ingassing** -> when carbon is subducted into the mantle of a planet

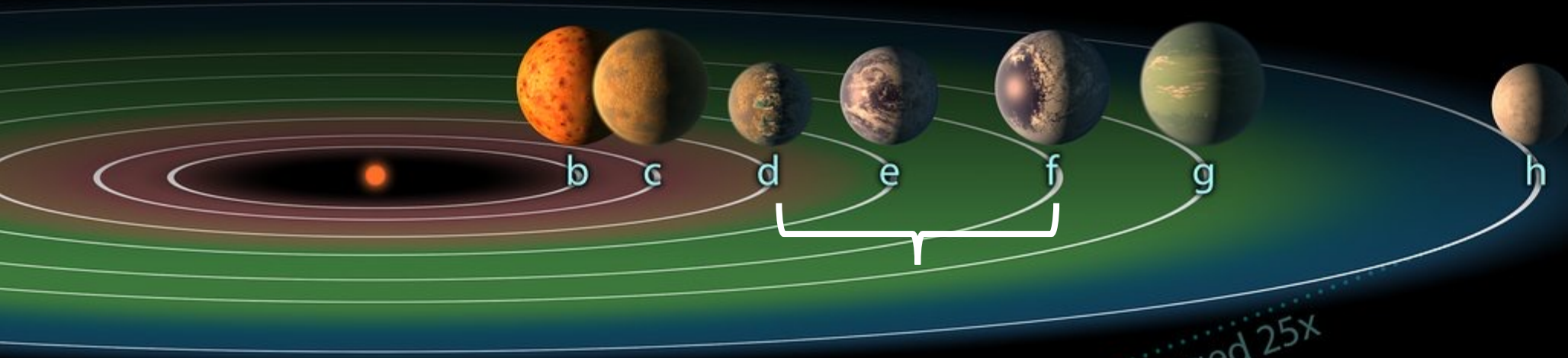
**Degassing or Outgassing** -> when carbon escapes from the mantle through volcanic eruptions and mid-ocean ridges

Subduction of carbon  
into the mantle

Adapted from Foley 2015

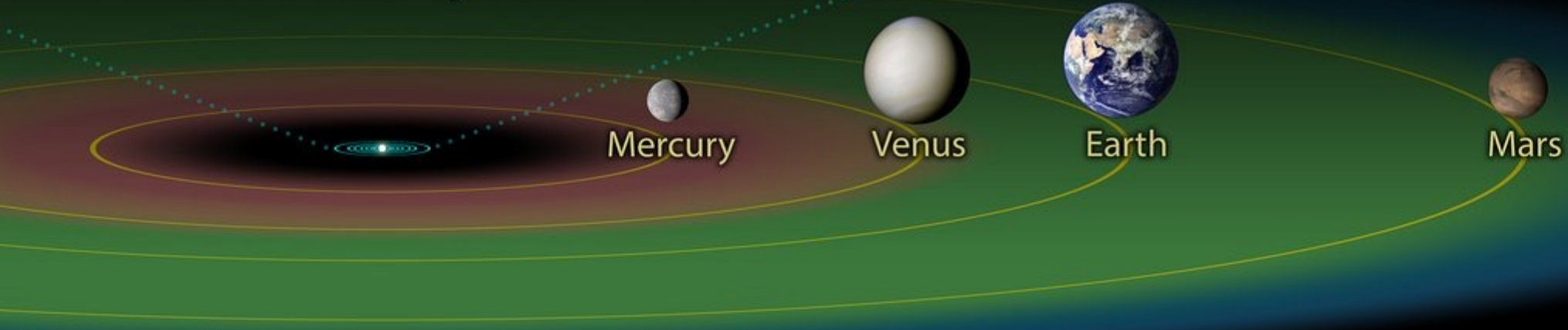
We mainly obtain atmosphere observations for exoplanets and therefore it is crucial to study their evolution and how interior processes can shape the atmosphere we observe today with telescopes!

## TRAPPIST-1 System



Enlarged 25x

## Inner Solar System

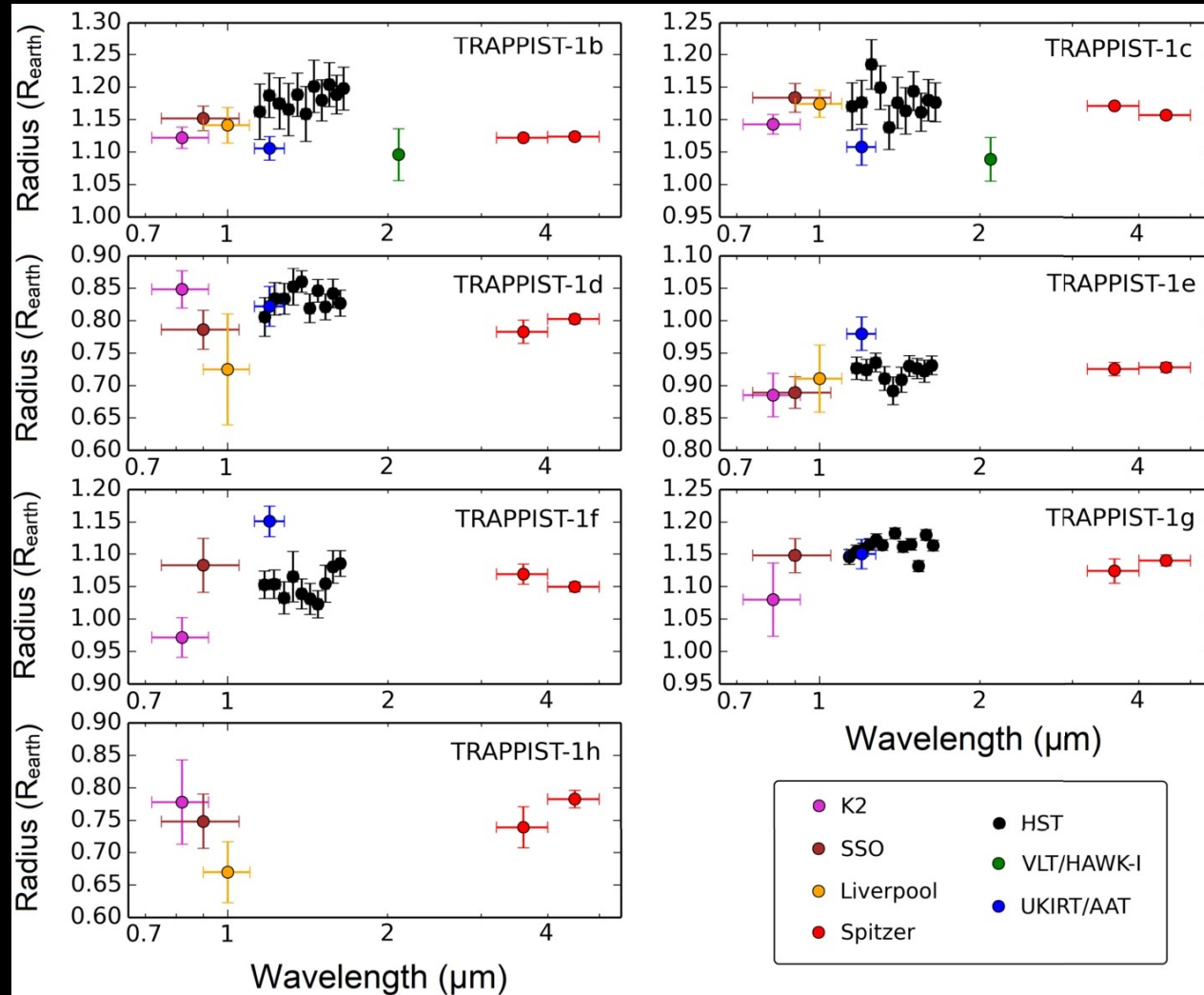


Illustration

Credit: NASA/JPL-Caltech



# The TRAPPIST-1 planets have been observed many times since their discovery



- It is very unlikely that the TRAPPIST-1 planets have hydrogen-dominated atmospheres
- It is expected that these planets either have a high molecular weight atmosphere or no atmosphere at all
- It is also very important to account for stellar contamination in observations!

Figure from Turbet et al. 2020

# TRAPPIST-1b was looked at with JWST!

Article | [Published: 27 March 2023](#)

## **Thermal Emission from the Earth-sized Exoplanet TRAPPIST-1 b using JWST**

[Thomas P. Greene](#) , [Taylor J. Bell](#), [Elsa Ducrot](#), [Achrène Dyrek](#), [Pierre-Olivier Lagage](#) & [Jonathan J. Fortney](#)

- They looked at the thermal emission from the planet with MIRI: found a dayside temperature of ~500K
- Suggests that it has no significant atmosphere!

*GOAL:*

To study the role of **outgassing** in the evolution of the atmospheric composition and planetary climate for TRAPPIST-1 d, e, f

How do we study this?





# Outgassing Models

We choose  
and test a  
range of  
oxygen  
fugacities  
( $fO_2$ )



We model the  
carbon cycle  
on the  
TRAPPIST-1  
planets



We track the  
following gas  
species:

- $H_2O$
- $H_2$
- $CO$
- $CO_2$

# Outgassing Models

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# Oxygen Fugacity ( $fO_2$ )

```
graph TD; A["Oxygen Fugacity (fO2)"] --> B["Controls mantle oxidation state"]; A --> C["Influences outgassing rates and composition"]; A --> D["Can be defined as a function of temperature also known as a redox buffer"];
```

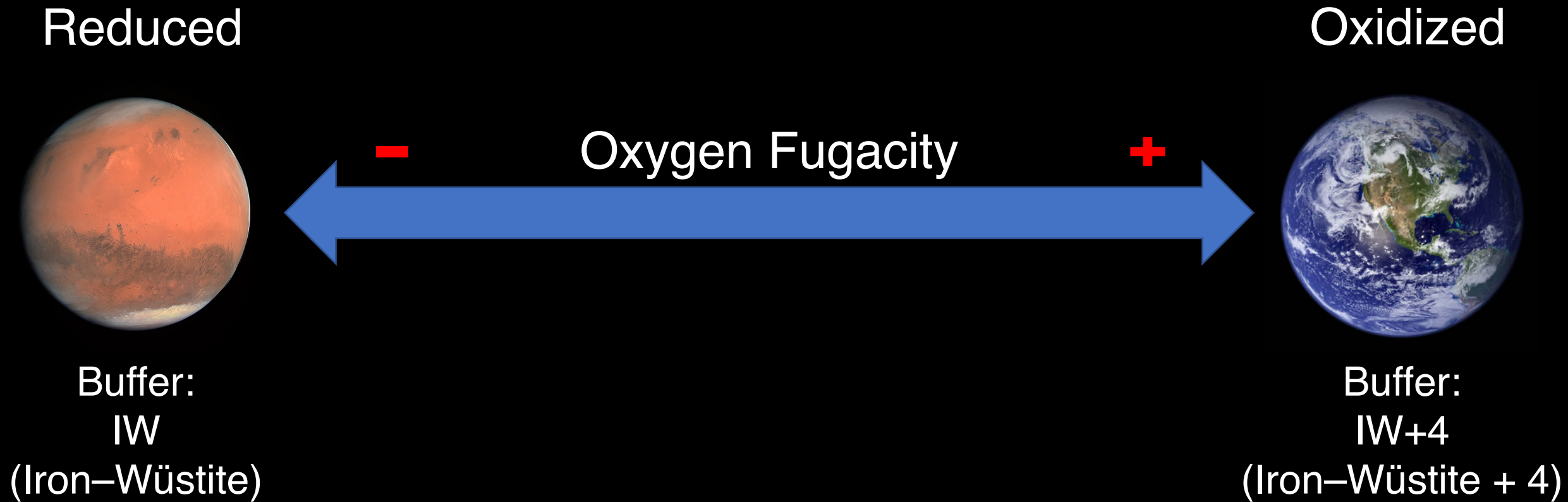
Controls  
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Influences  
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Can be defined  
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a ***redox buffer***



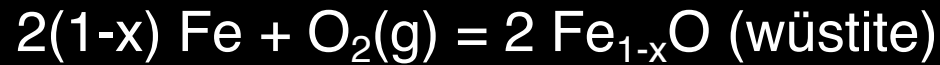
We consider a range of oxygen fugacities ( $f\text{O}_2$ ) for TRAPPIST-1 d, e and f



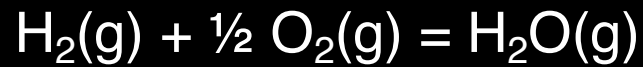


# Oxygen Fugacity ( $f_{O_2}$ )

Iron – Wüstite (IW) buffer



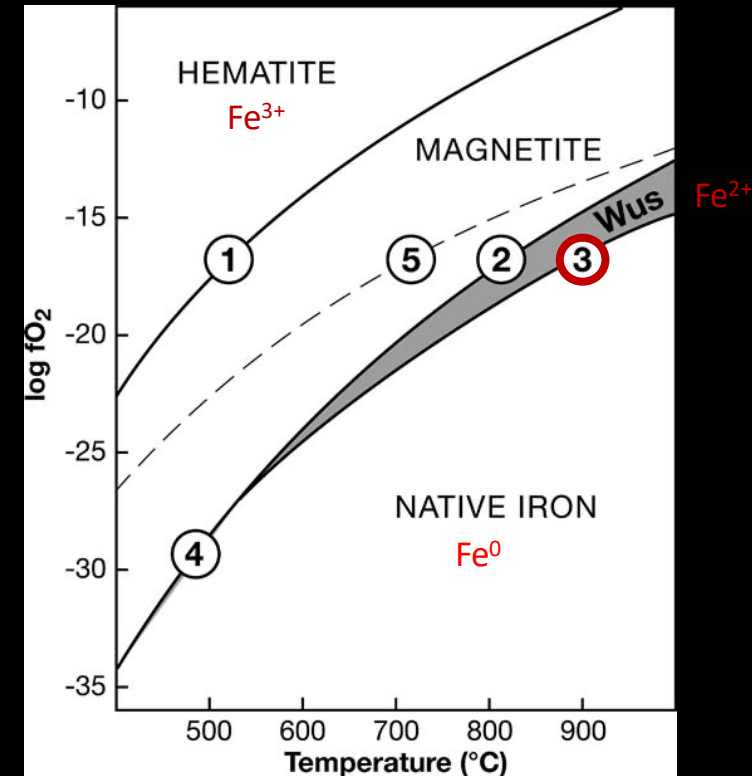
Gas reactions



$$\frac{P_{\text{H}_2\text{O}}}{P_{\text{H}_2}} \propto f_{\text{O}_2}^{1/2}$$



$$\frac{P_{\text{CO}_2}}{P_{\text{CO}}} \propto f_{\text{O}_2}^{1/2}$$

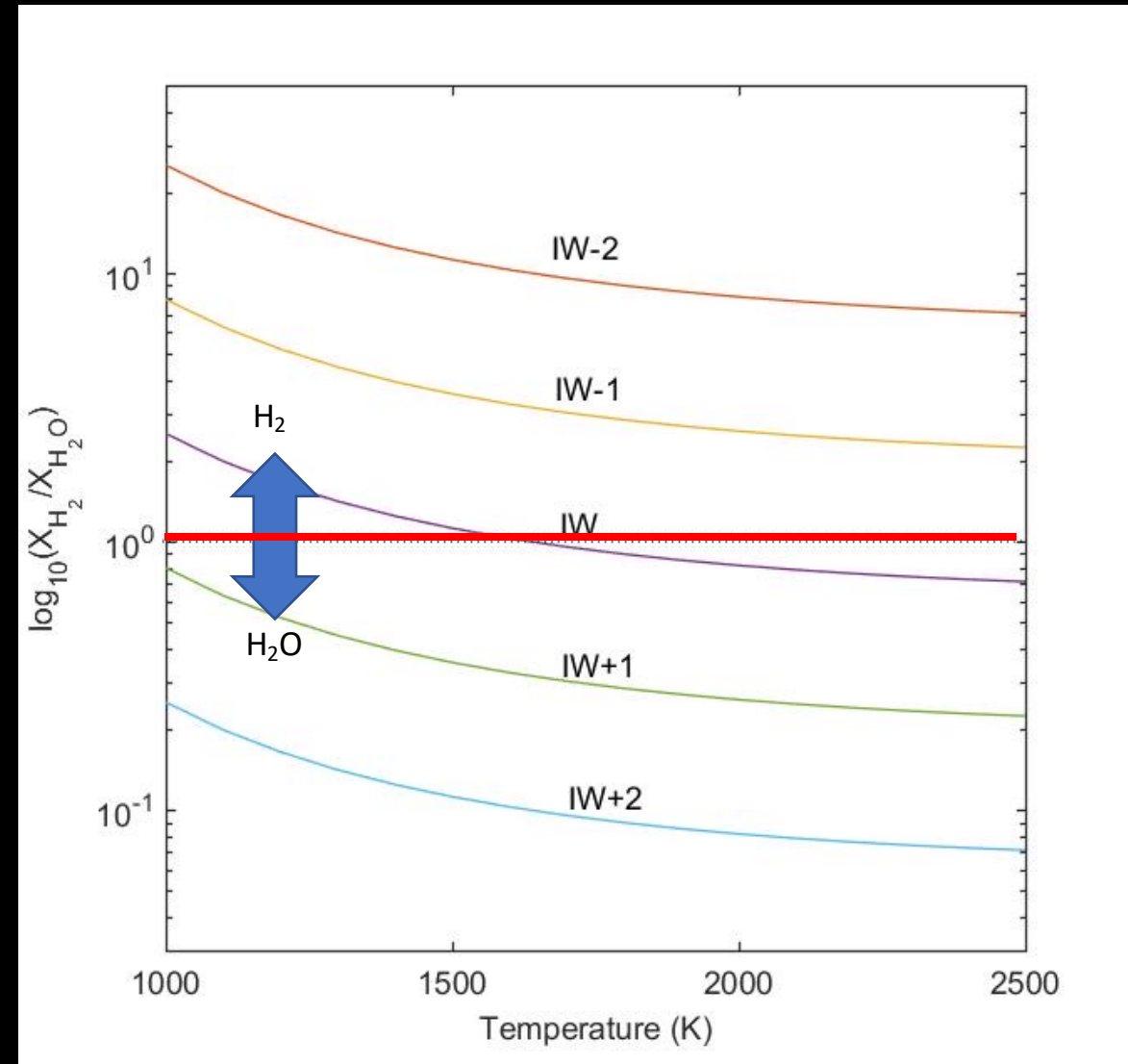
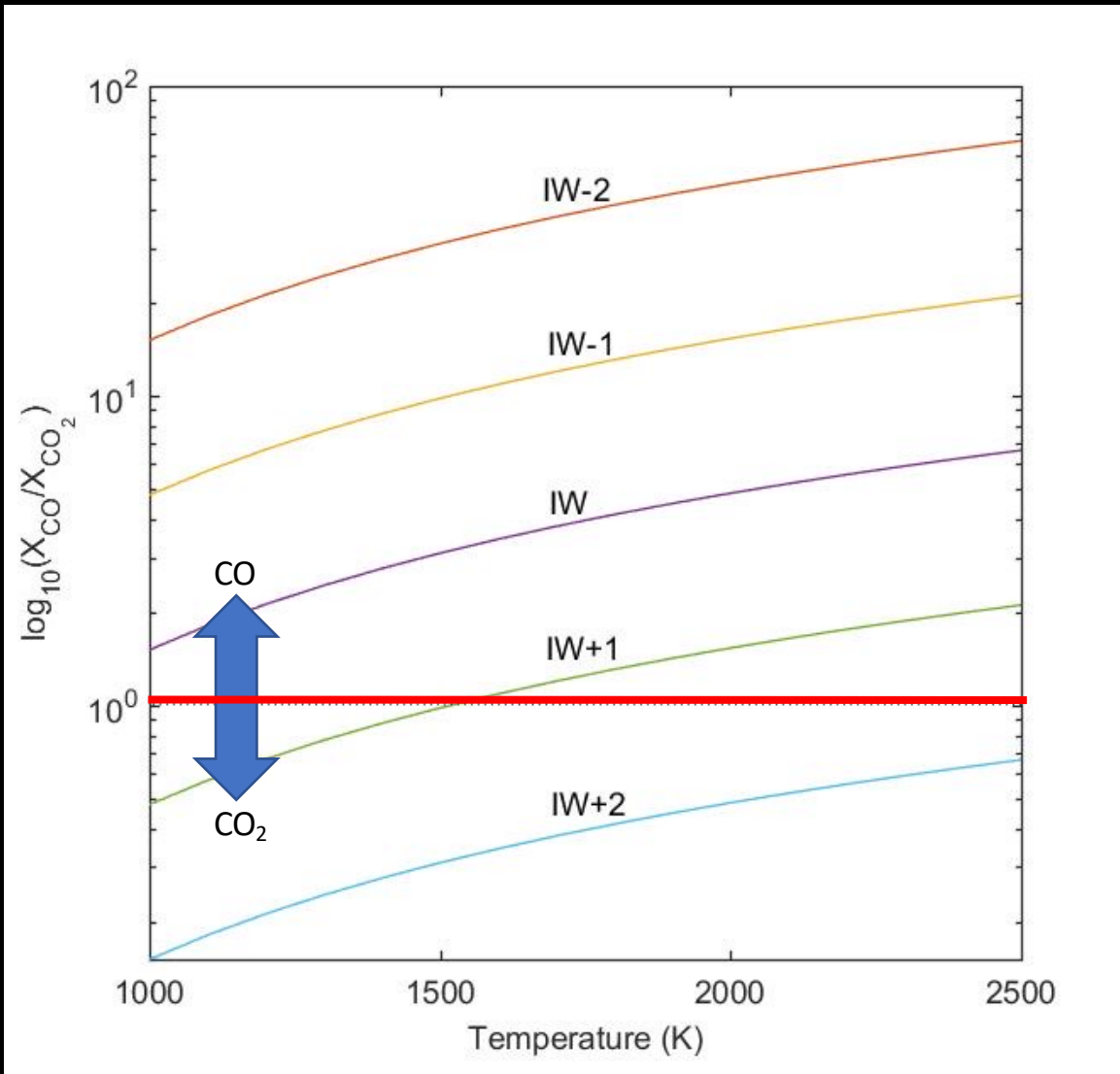


Seifert et al. (2010) Eu. J. Min.

- Controls relative gas abundances (e.g.  $\text{H}_2/\text{H}_2\text{O}$ ,  $\text{CO}/\text{CO}_2$ ).
- Carbon outgassing rates depend on it.

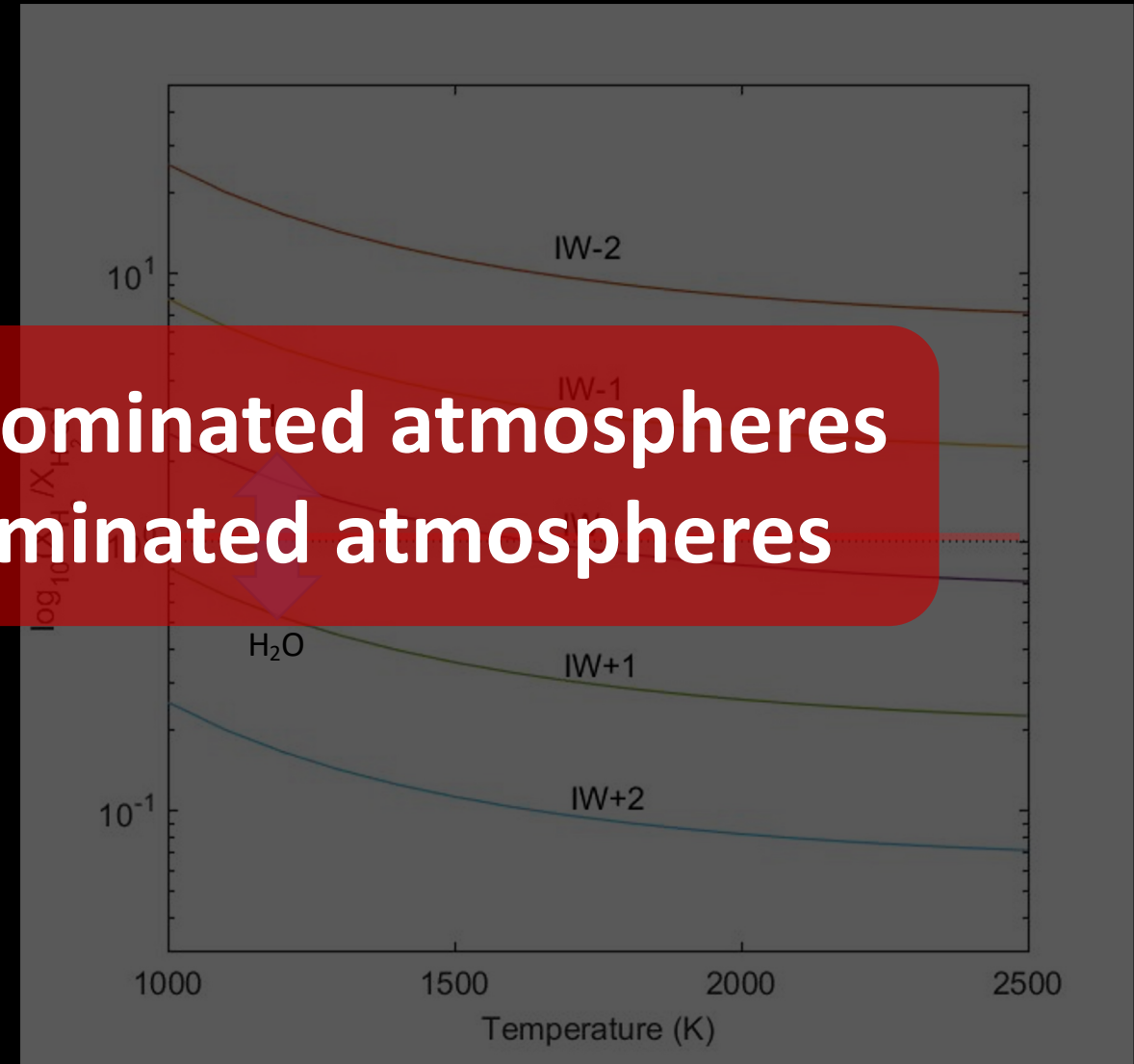
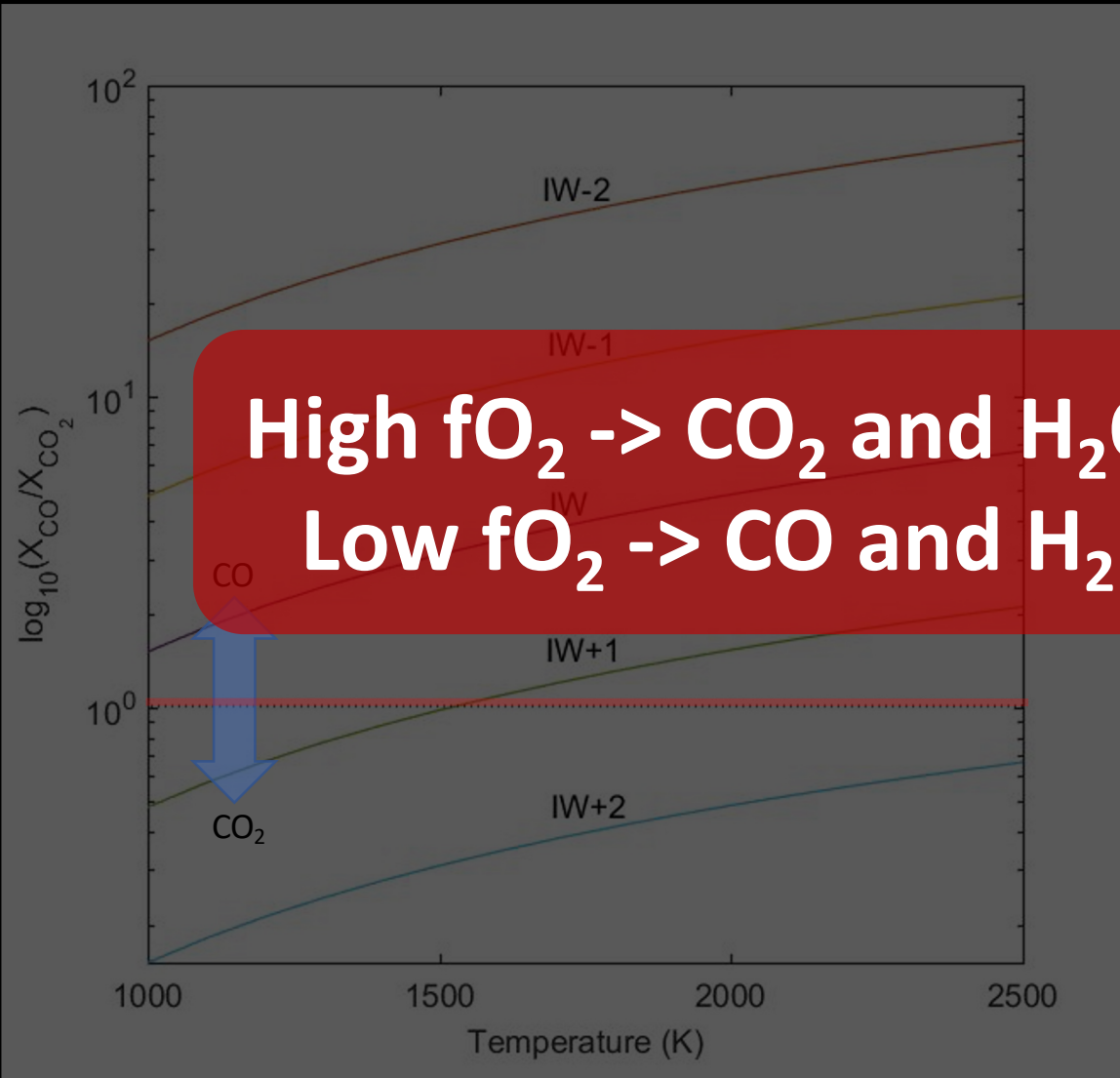
Modified from: Laura Schaefer

# Gas Composition at fixed $fO_2$



# Gas Composition at fixed $fO_2$

High  $fO_2 \rightarrow CO_2$  and  $H_2O$  dominated atmospheres  
Low  $fO_2 \rightarrow CO$  and  $H_2$  dominated atmospheres



# Outgassing Models

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We model the  
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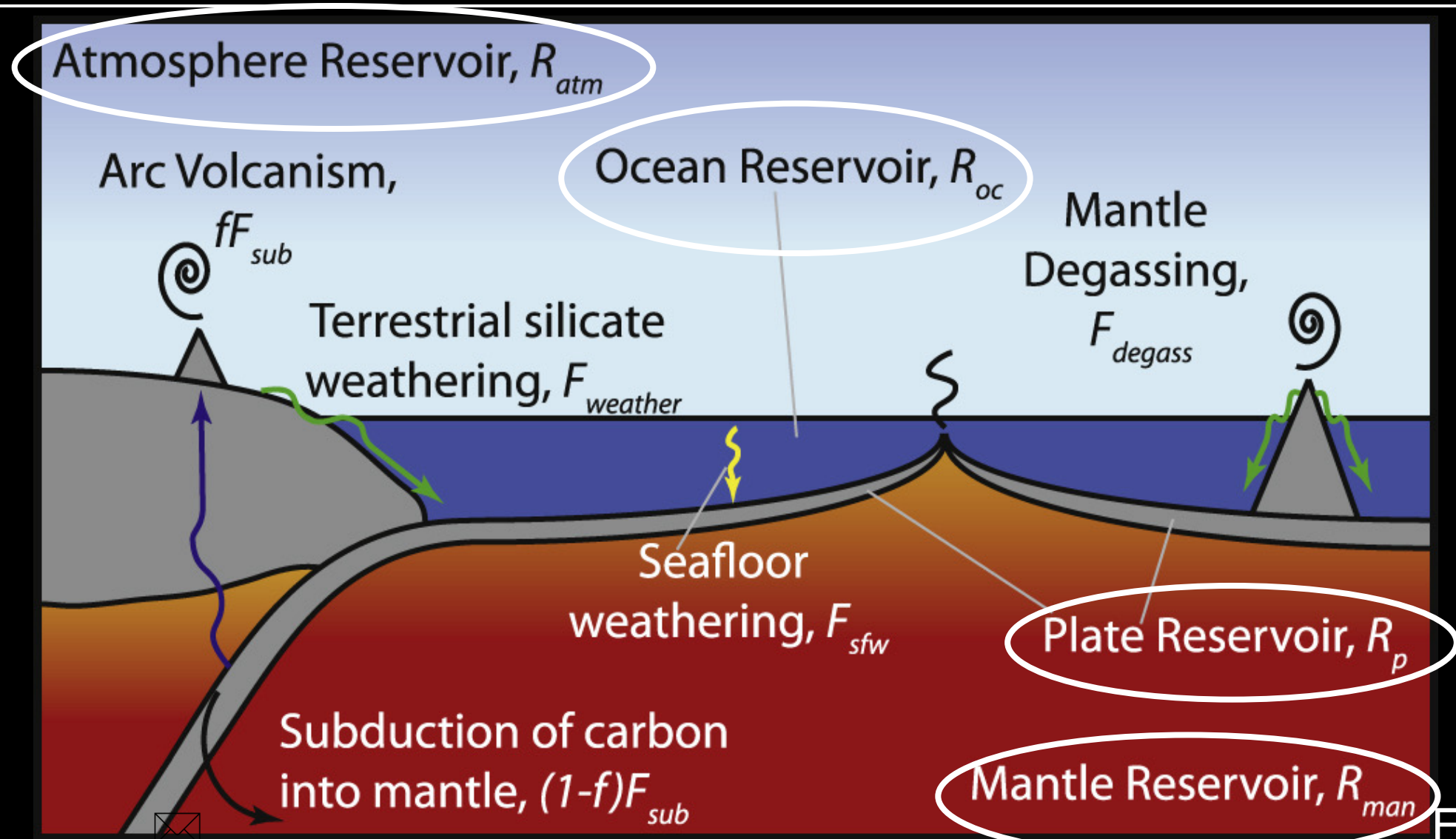


We track the  
following gas  
species:

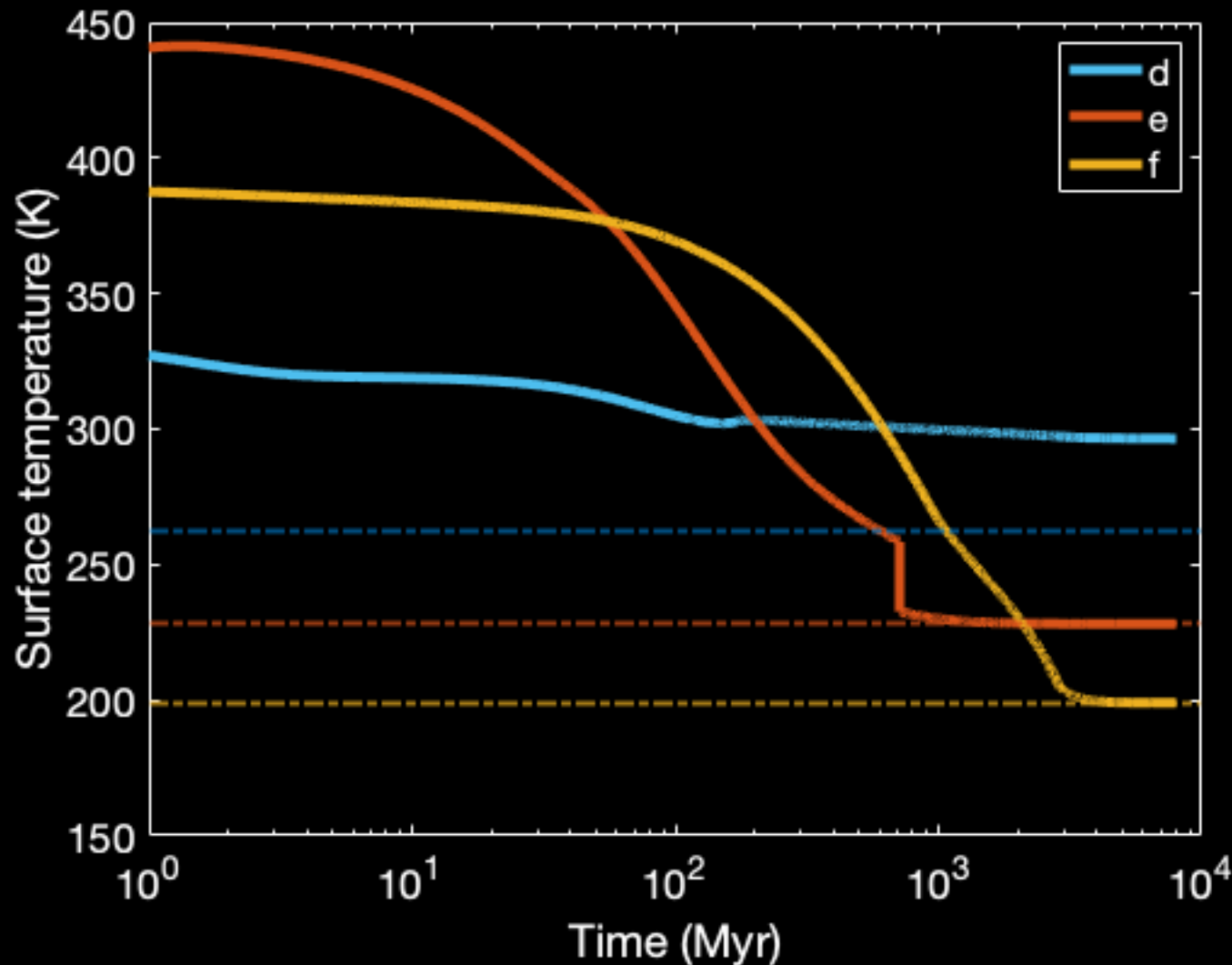
- $H_2O$
- $H_2$
- $CO$
- $CO_2$



We model the carbon cycle and incorporate the deep water cycle (Schaefer & Sasselo 2015) to track the gas species



# We use Driscoll & Bercovici 2013 and Foley 2015 to calculate **surface temperatures** for TRAPPIST-1 planets



Driscoll & Bercovici 2013:

$$T_s^4 = T_e^4 \left( 1 + \frac{3}{4} \tau_g \right) \text{ where:}$$

$$\tau_g = \frac{P_w}{P_{k,w}} + \frac{P_c}{P_{k,c}}$$

$T_s$  = surface temperature

$T_e$  = equilibrium temperature

$\tau_g$  = optical depth at surface

$P_w$  = water partial pressure

$P_{k,w}$  = water opacity pressure

$P_c$  = carbon dioxide partial pressure

$P_{k,c}$  = carbon dioxide opacity pressure

Foley 2015:

$$T = T^* + 2 (T_e - T_e^*) + 4.6 \left( \frac{P_c}{P_c^*} \right)^{0.346} - 4.6$$

Our model also accounts for photochemistry and atmospheric escape

Outgassing Model

```
graph TD; A[Outgassing Model] --> B[Photochemistry: CO + OH -> CO2 + H]; A --> C[Hydrogen escape]
```

Photochemistry:  
 $CO + OH \rightarrow CO_2 + H$

Hydrogen escape

# Outgassing Models

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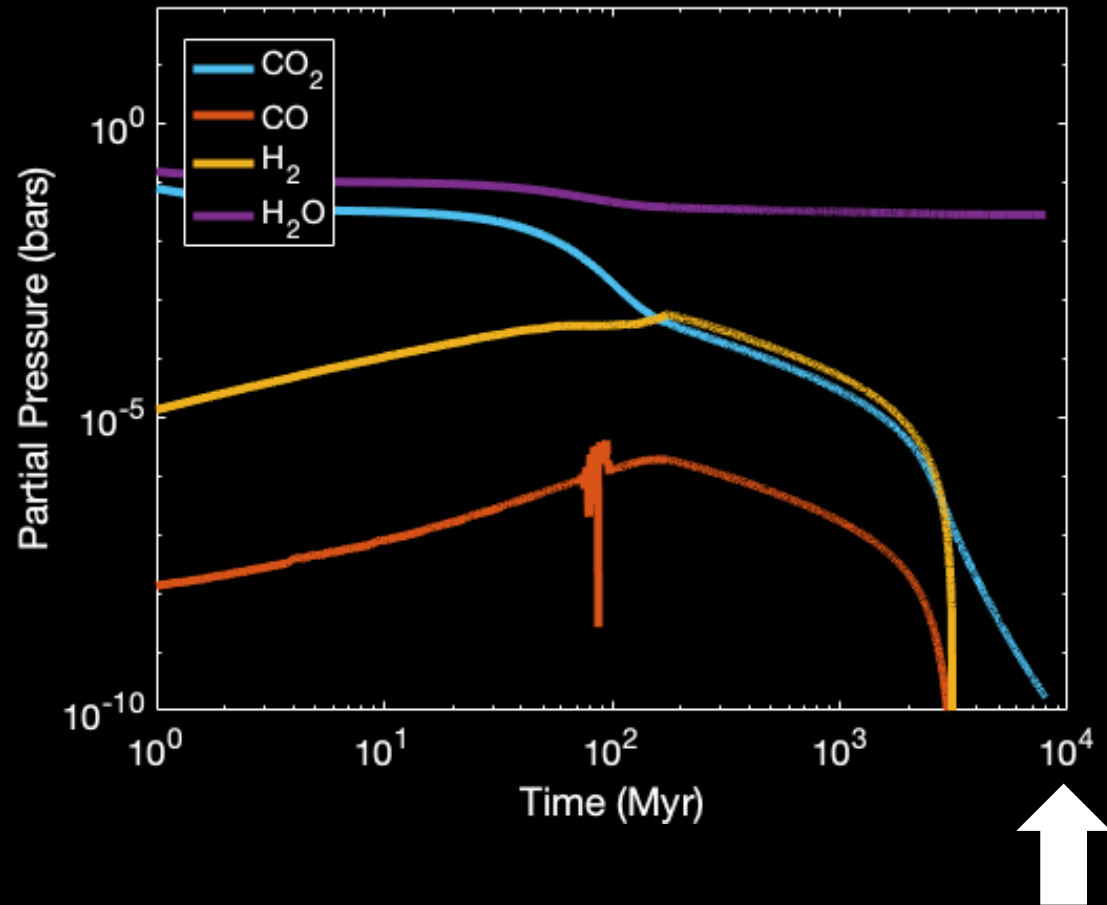
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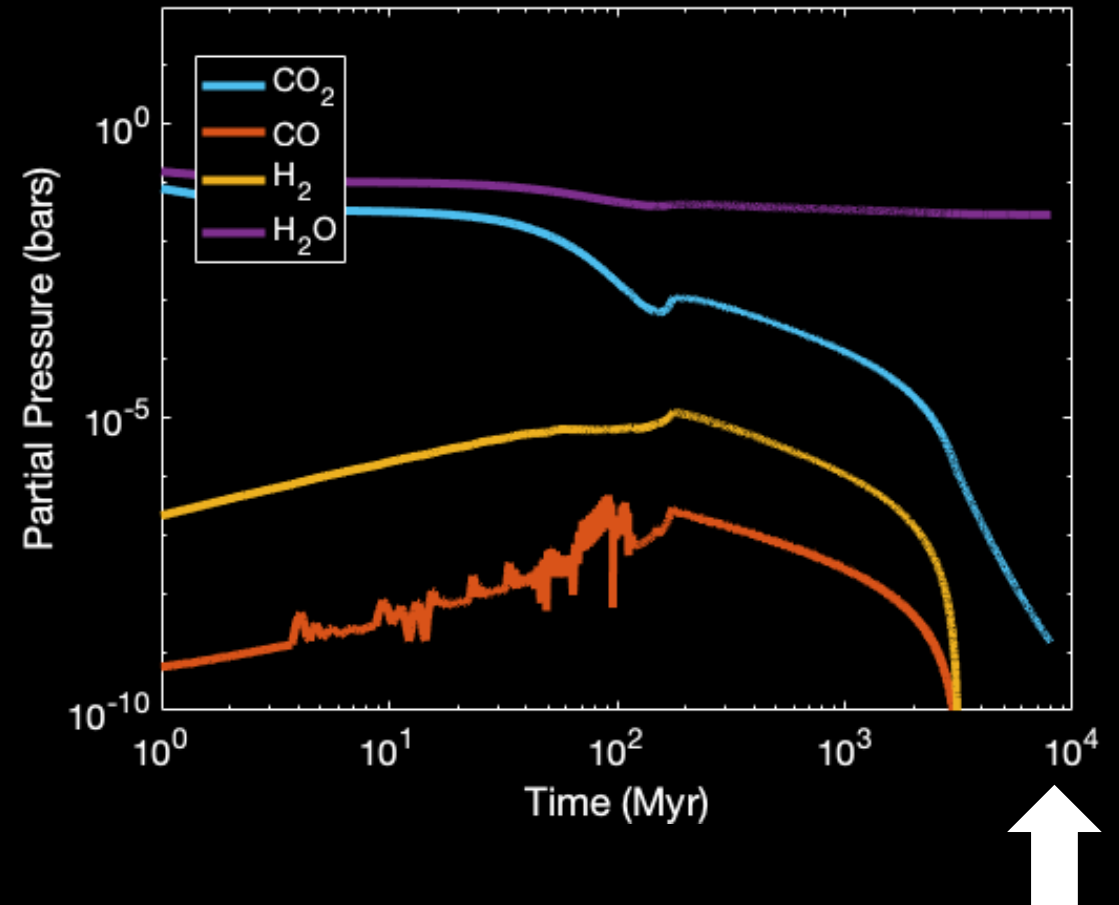


# Atmospheric Composition of TRAPPIST-1 d

Low  $f\text{O}_2$  (IW)



High  $f\text{O}_2$  (IW + 4)

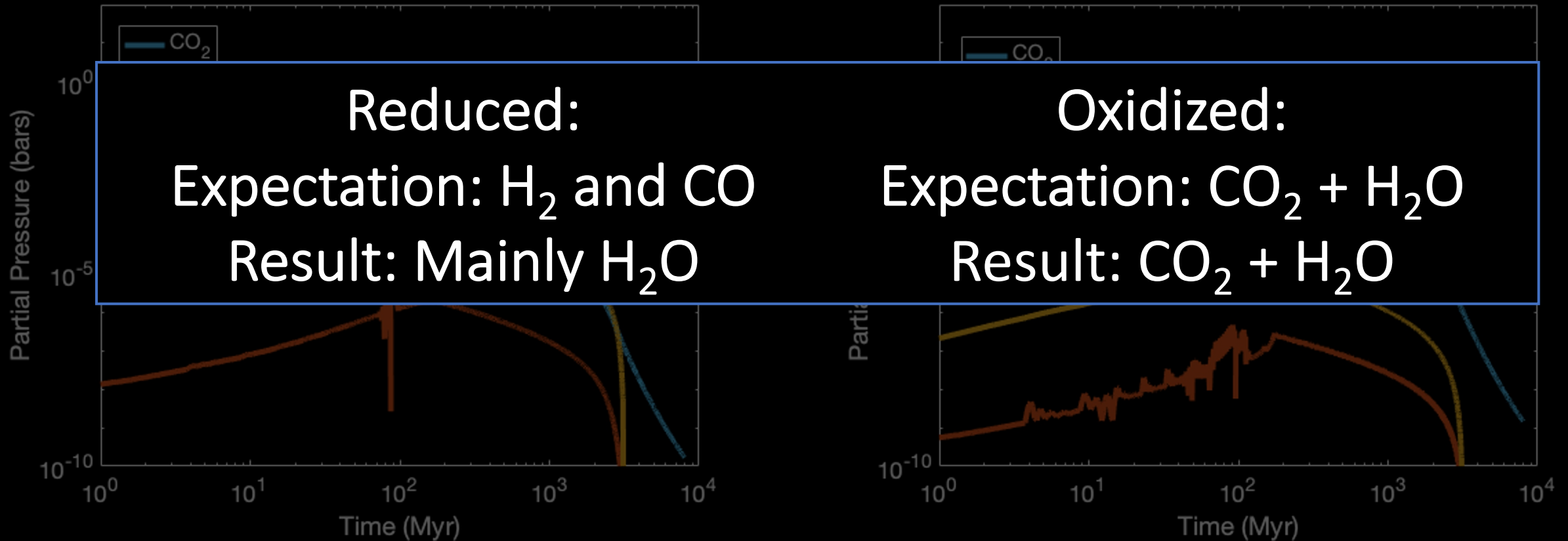


Gonzalez-Quiles *et al.* 2023 (In Prep)

# Atmospheric Composition of TRAPPIST-1 d

Low  $f\text{O}_2$  (IW)

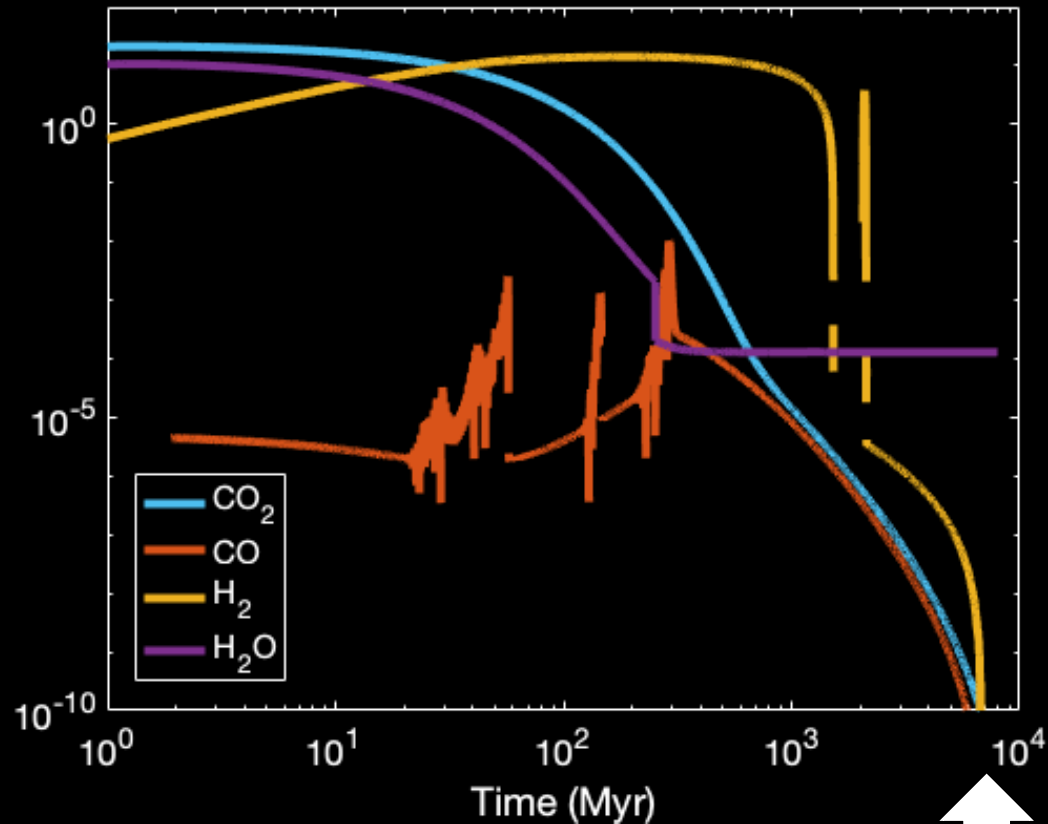
High  $f\text{O}_2$  (IW + 4)



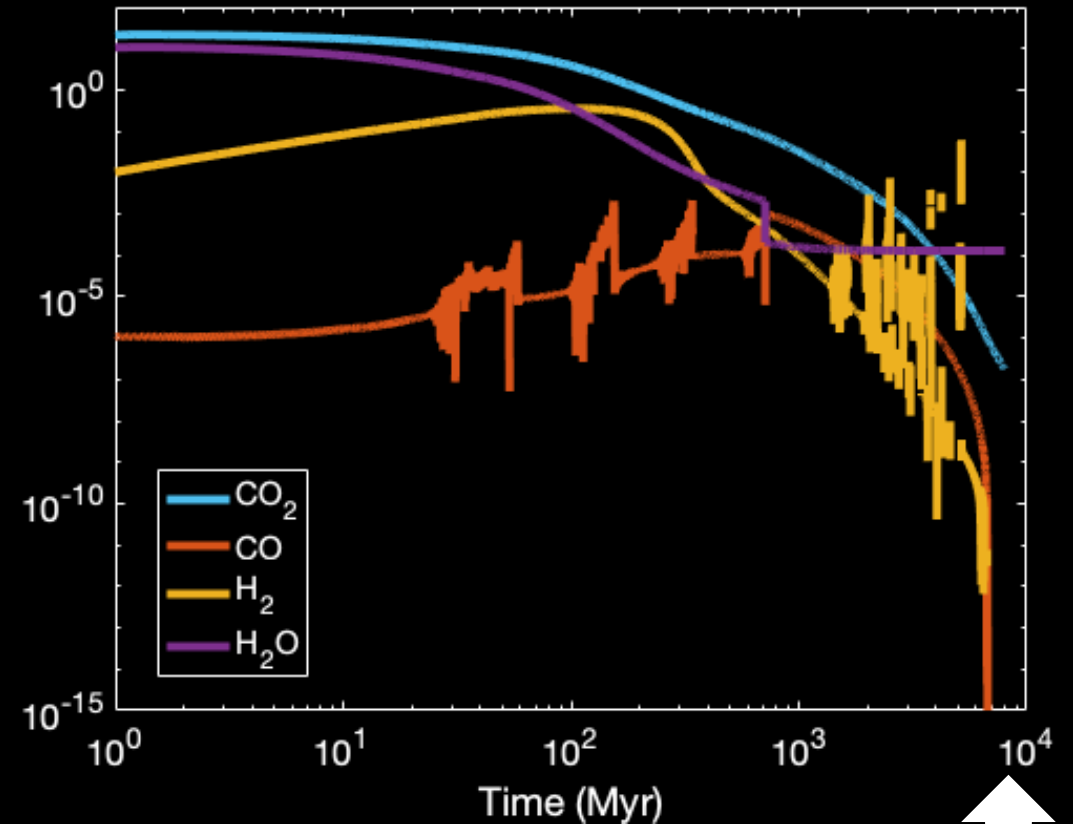
Gonzalez-Quiles *et al.* (In Prep)

# Atmospheric Composition of TRAPPIST-1 e

Low  $f\text{O}_2$  (IW)



High  $f\text{O}_2$  (IW + 4)

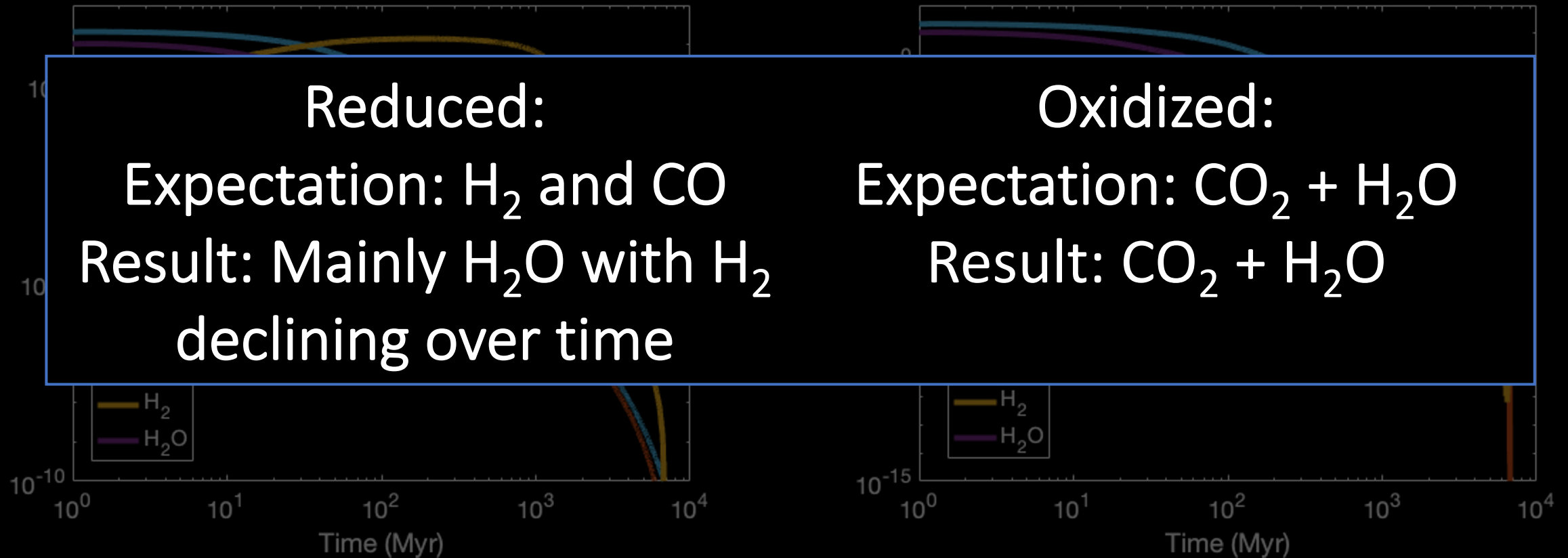


Gonzalez-Quiles *et al.* 2023 (In Prep)

# Atmospheric Composition of TRAPPIST-1 e

Low  $f\text{O}_2$  (IW)

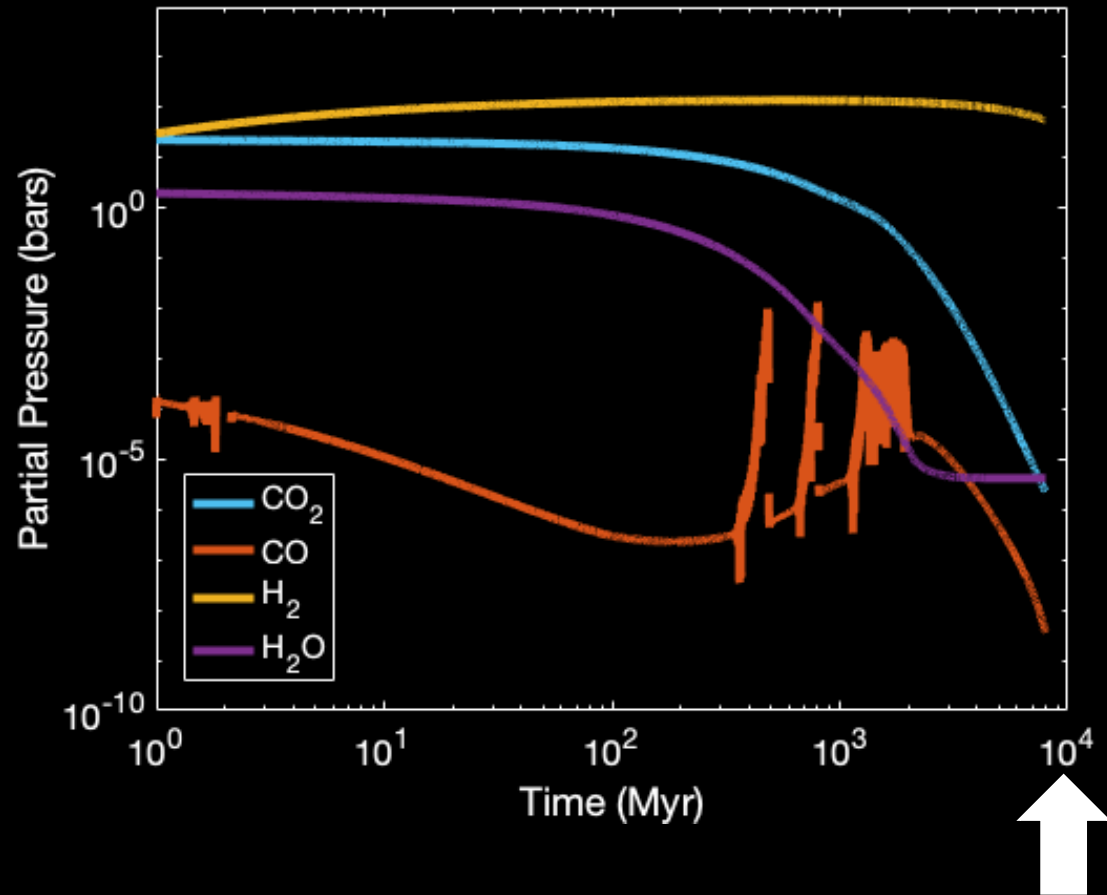
High  $f\text{O}_2$  (IW + 4)



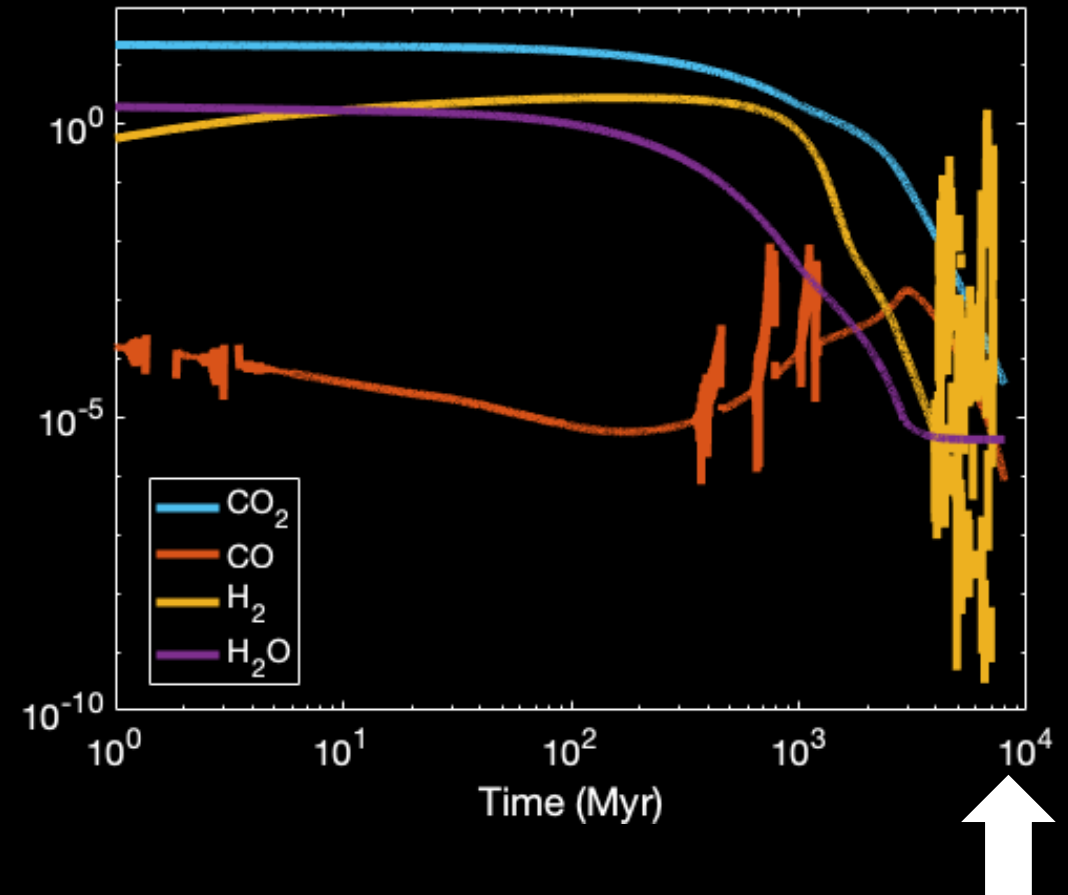


# Atmospheric Composition of TRAPPIST-1 f

Low  $f\text{O}_2$  (IW)



High  $f\text{O}_2$  (IW + 4)

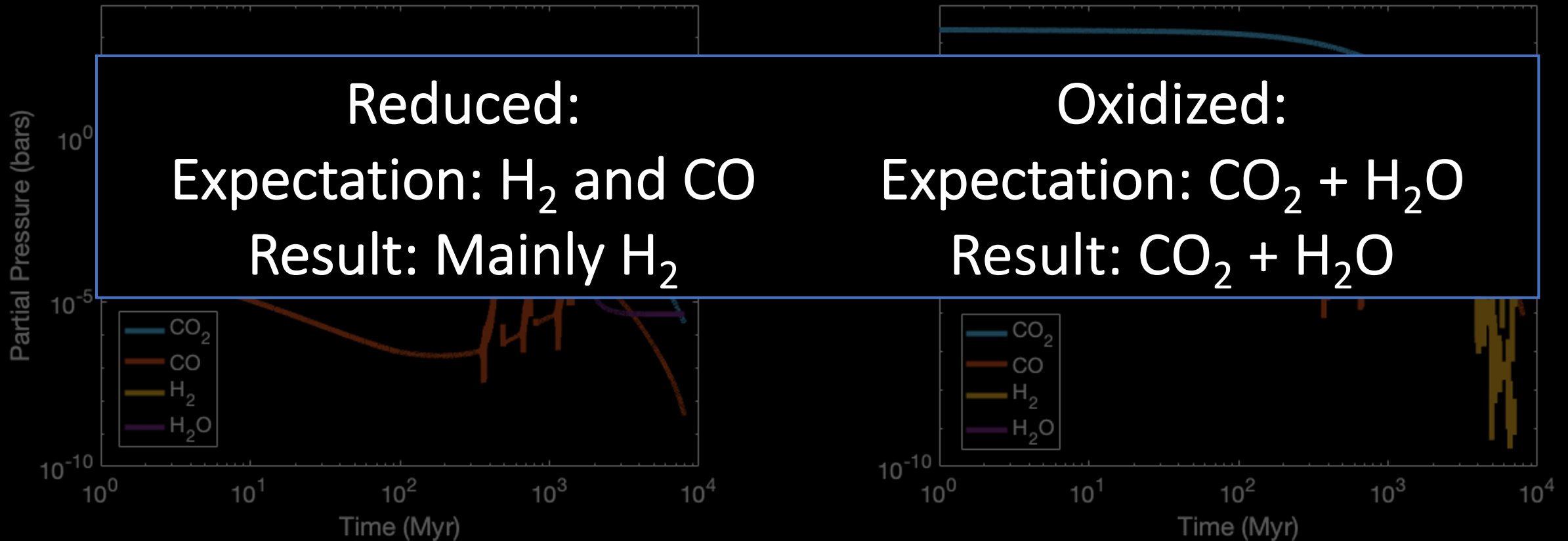


Gonzalez-Quiles *et al.* 2023 (In Prep)

# Atmospheric Composition of TRAPPIST-1 f

Low  $f\text{O}_2$  (IW)

High  $f\text{O}_2$  (IW + 4)



Our results suggest that the atmospheric composition of these planets may depend on the oxygen fugacity and outgassing rates, but other important processes can also impact the atmospheric composition of these planets!

# What's Up Next?

Add nitrogen and sulfur into our models.

Modify the thermal evolution model to incorporate heat transport from the core to the mantle.

Interested in collaborating, please reach out to me!

Website:



jgonza70@jhu.edu



@JunellieG







## Summary

Modeling the deep water cycle and the carbon cycle is important for understanding interior-surface-atmosphere interactions and their effect on habitability of exoplanets.

Outgassing models were run for TRAPPIST-1 d, e and f for different oxygen fugacities.

Oxygen fugacity can impact outgassing and therefore influence the atmospheric composition of planets.

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- Prof. Sabine Stanley and their research group
- Carnegie EPL Astro
- LUMA (League of Underrepresented Minoritized Astronomers)
- Tyler Perez
- All my friends
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  - PI: Kevin Stevenson, Sci-PI: Ravi Kopparapu
  - Grant No. 80NSSC21K0905
- This work is also supported by the NSF GRFP
  - JHU award No. DGE-1746891

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# TRAPPIST-1 Parameters

TRAPPIST -1 Planets	Radius ( $R_{\text{earth}}$ )	Mass ( $M_{\text{earth}}$ )	$g$ ( $g_{\text{Earth}}$ )	$T_{\text{eq}}$ (K)	Core Mass Fraction (wt %)	Water Mass Fraction (wt %)	$R_{\text{core}}$ ( $R_{\text{earth}}$ )
d	0.788	0.388	0.624	262	19.7	$10^{-3}$ ( $10^{-4}$ , 2)	0.38
e	0.92	0.692	0.817	228	24.6	0.3 ( $10^{-4}$ , 2.1)	0.48
f	1.045	1.039	0.951	199	20.1	1.0 (0.6, 3.4)	0.49

Agol et al. 2021 PSJ 2:1